

# Out of Flatland: 3D Transport Barriers in Oceanography

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# Lagrangian Coherent Structures

- Defined by Haller (2000) in terms of FTLE (default metric used here)
- Other LCS characterizations
  - Joseph & Legras (2002) - FSLE
  - Mancho & Mendoza (2010) - minimal trajectories
  - Rypina - complexity
  - Mezic et al (2010) - mesohyperbolicity
  - Haller (2011) - geodesic material surfaces
- Most studies in GFD confined to 2D velocities
- Yet theory applies to  $\mathbf{R}^n$



# Are LCS important in GFD?

- MODE/POLYMODE (circa 1975) - Mesoscale eddies transport heat, salinity, and momentum

*But*

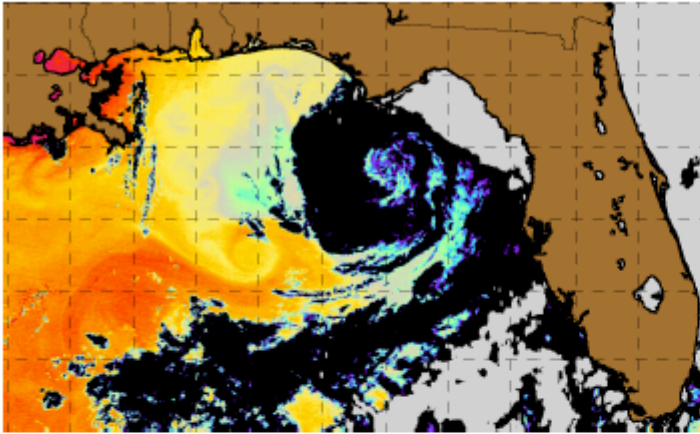
- How do eddies form?
- How many eddies are there?
- How do eddies exchange heat, etc with environment?
- Since MODE/POLYMODE
  - Growing Lagrangian user community
  - Dramatic oil spills
- Circa 1990 – Little Compton meeting. DST methods applied to 2D mesoscale and submesoscale transport

But Transport is *Volume* per time

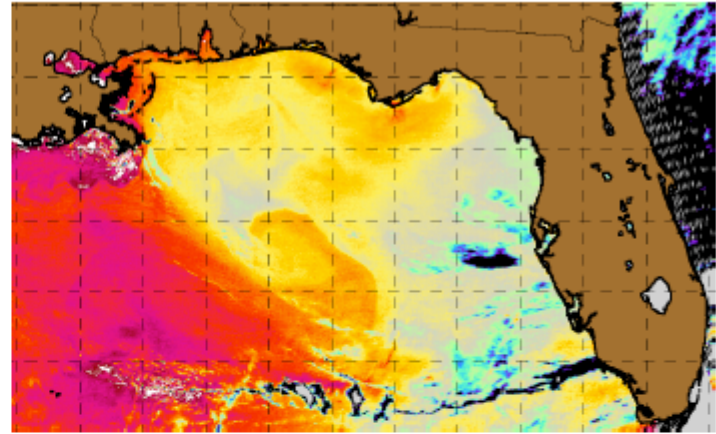
- 3D computations too costly (Garth et al 2007)
- 3D velocities not always available in GFD  
(Branicki & Kirwan 2010, Bettencourt et al 2012)
- *Question:*  
Can 3D LCS be estimated from 2D velocities?

# Are LCS Important in Oceanography?

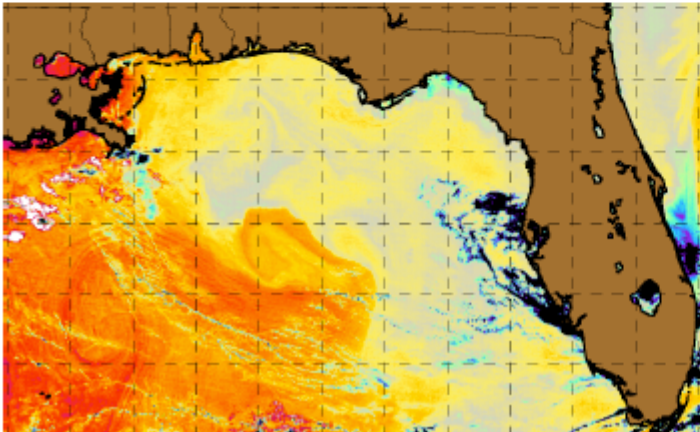
June 26



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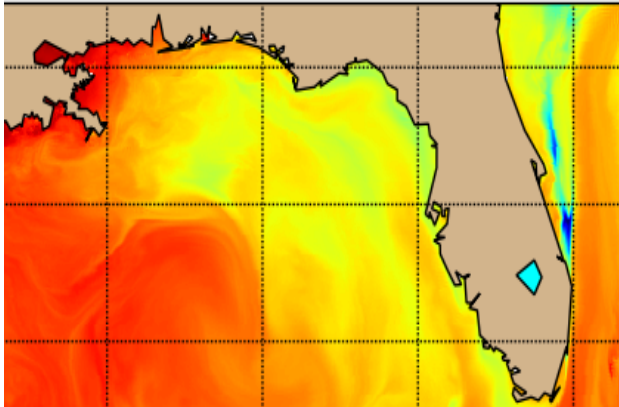


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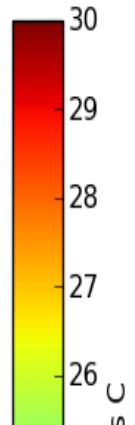
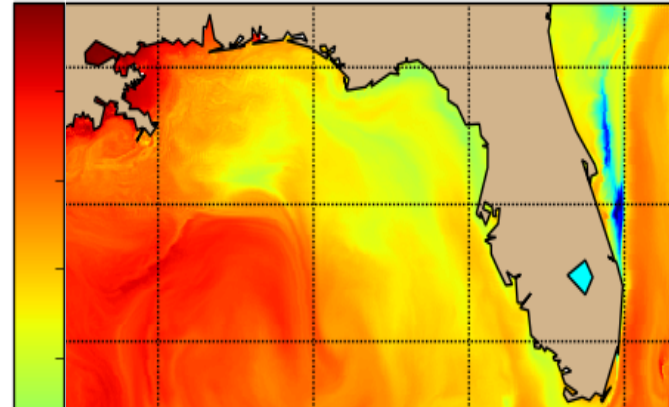


# Are LCS Important in Oceanography?

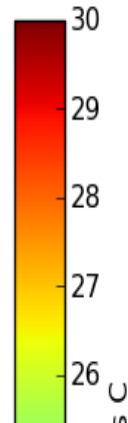
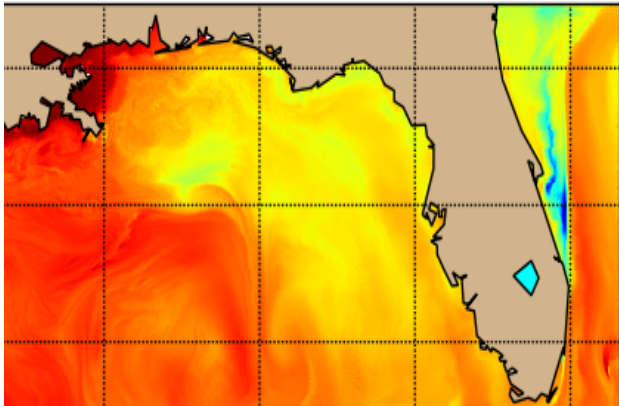
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# Out of Flatland – What if there was a 3<sup>rd</sup> Dimension?

- Mezic and Wiggins (1994)
- Toy Problem

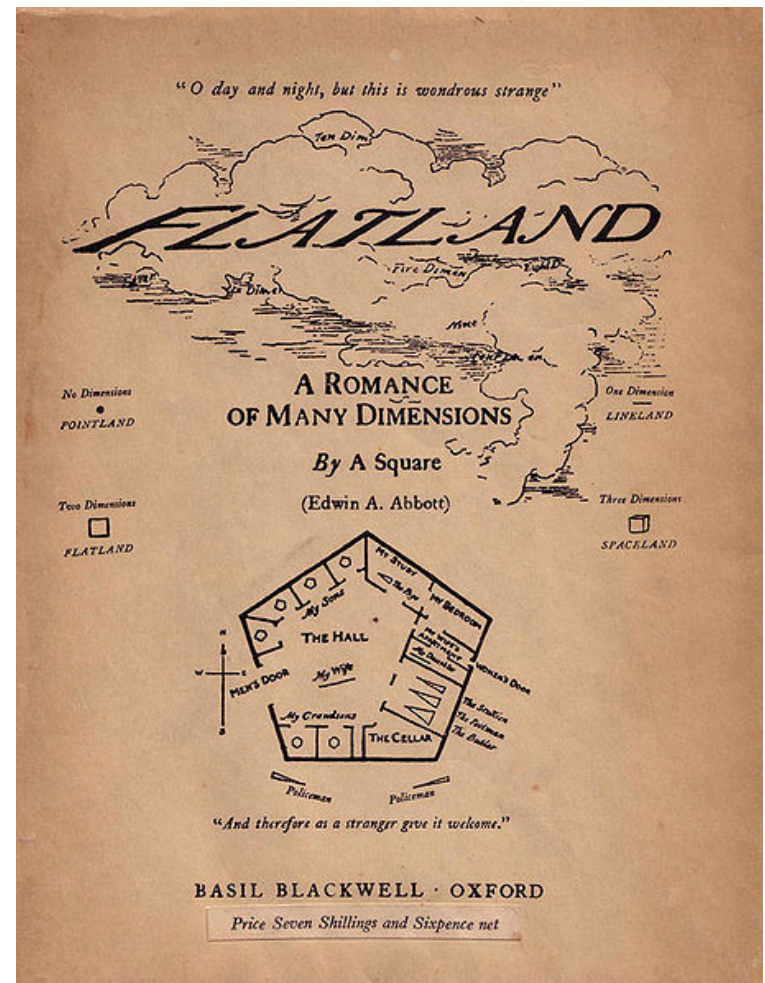
$$\dot{x} = \lambda_1 x - \omega y$$

$$\dot{y} = \omega x + \lambda_2 y$$

$$\dot{z} = -(\lambda_1 + \lambda_2) z$$

$$0 < \omega^2 - \lambda^2$$

$$\lambda = \frac{\lambda_1 + \lambda_2}{2}$$



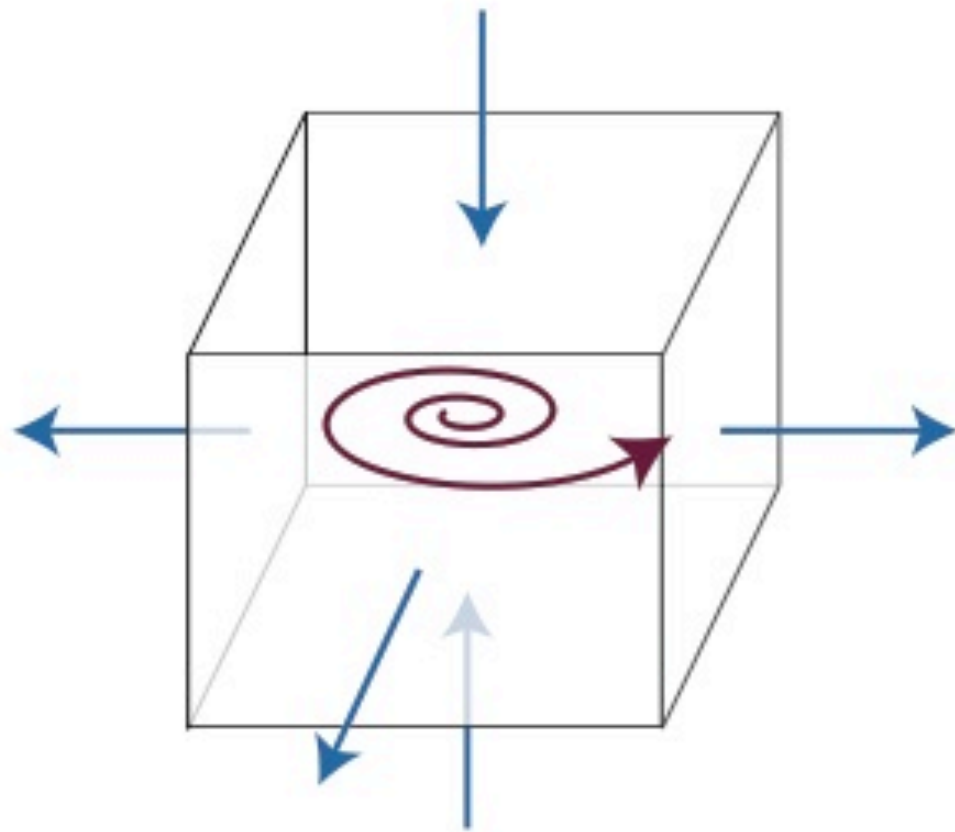
# Out of Flatland Solution

$$x = \exp \lambda t [X \cos \Omega t - (\lambda_2 X + \omega Y) \sin \Omega t / \Omega]$$

$$y = \exp \lambda t [Y \cos \Omega t - (\lambda_1 Y - \omega X) \sin \Omega t / \Omega]$$

$$z = Z \exp -2\lambda t$$

$$\Omega^2 = \omega^2 - \lambda^2$$

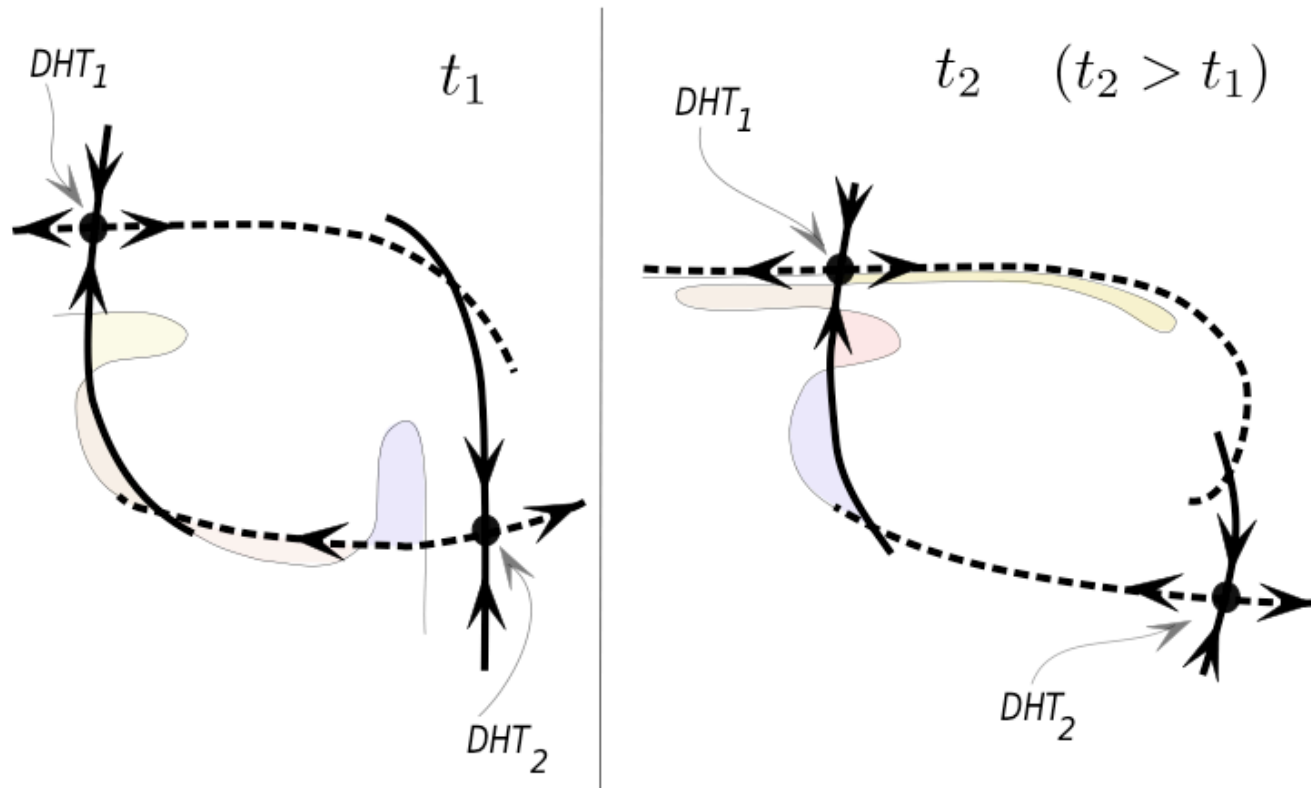


# Branicki & Kirwan (IJES 2010)

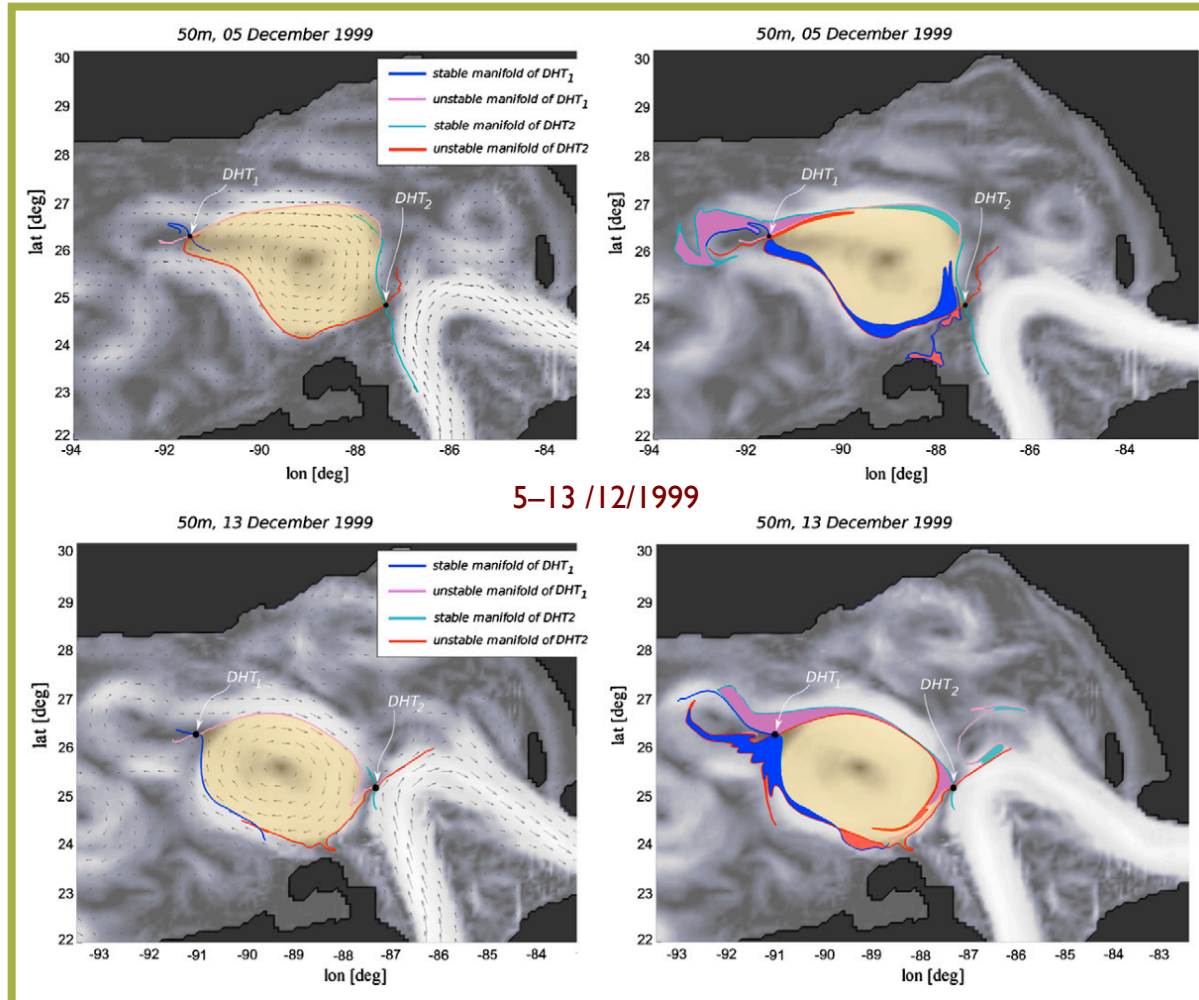
- FTLE located DHTs
- 1D manifolds using 2D velocities (Ide et al, 2002) from 0 to 250m for eddy Juggernaut
- Stitched 1D manifolds into 2D material surfaces



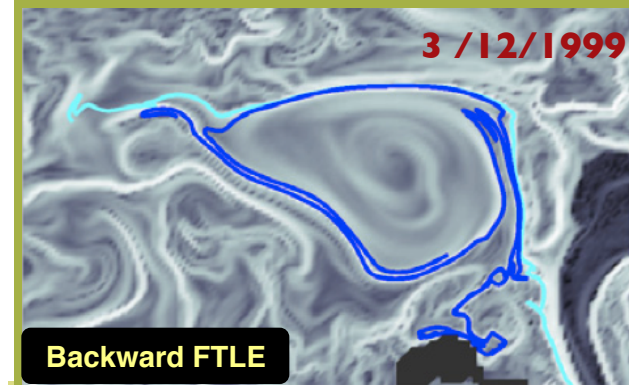
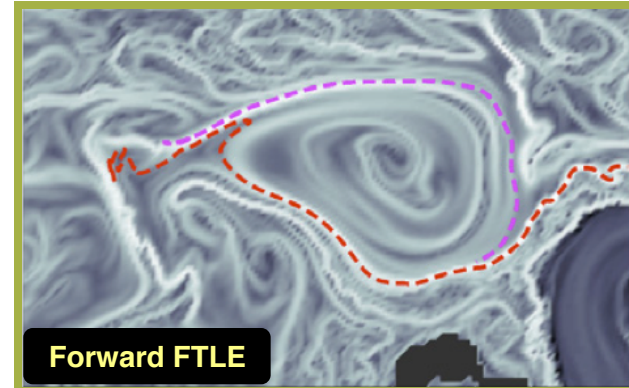
# Dynamical Systems Approach



# Loop Current Ring Exchange



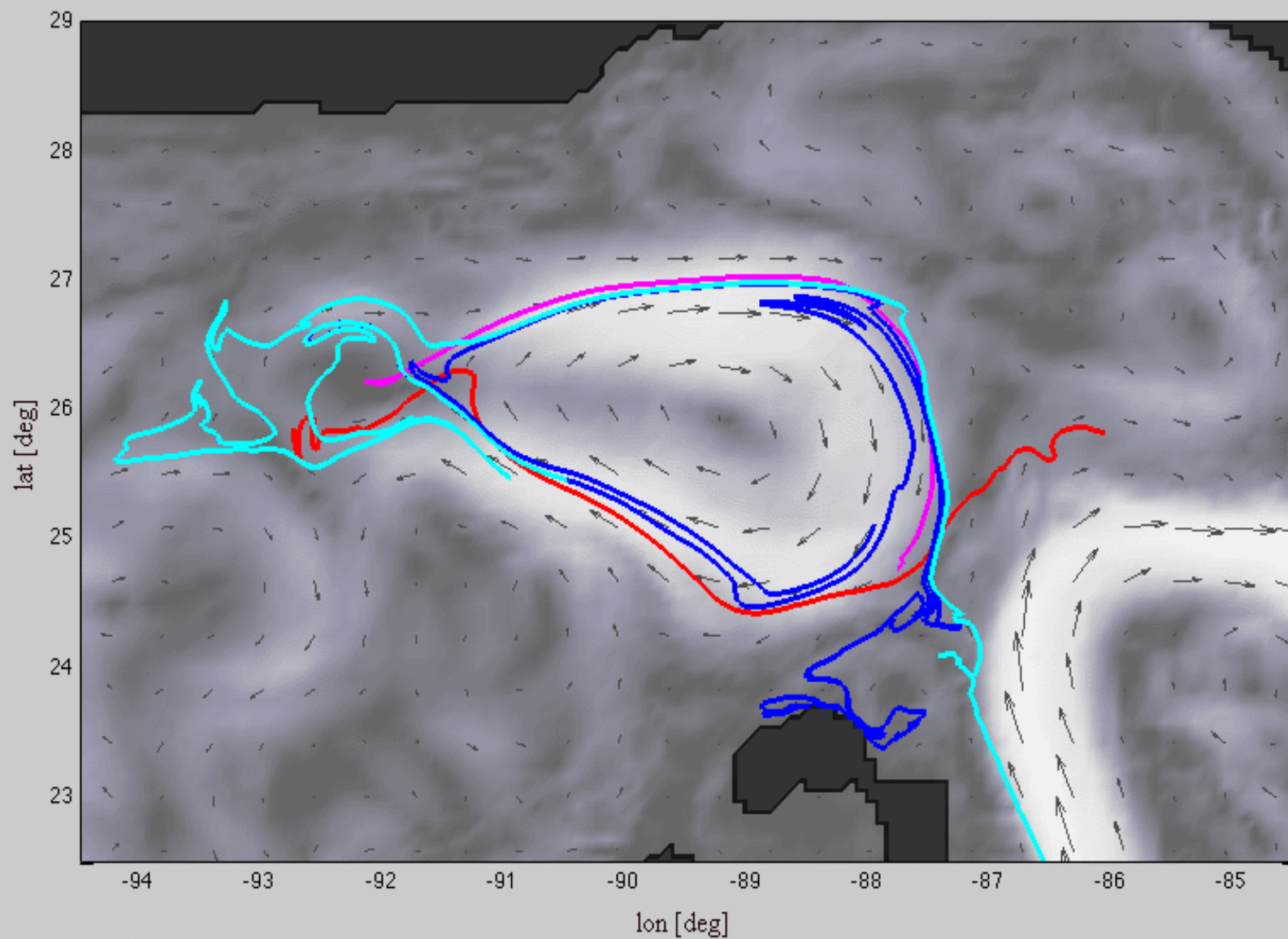
## Eddy Juggernaut

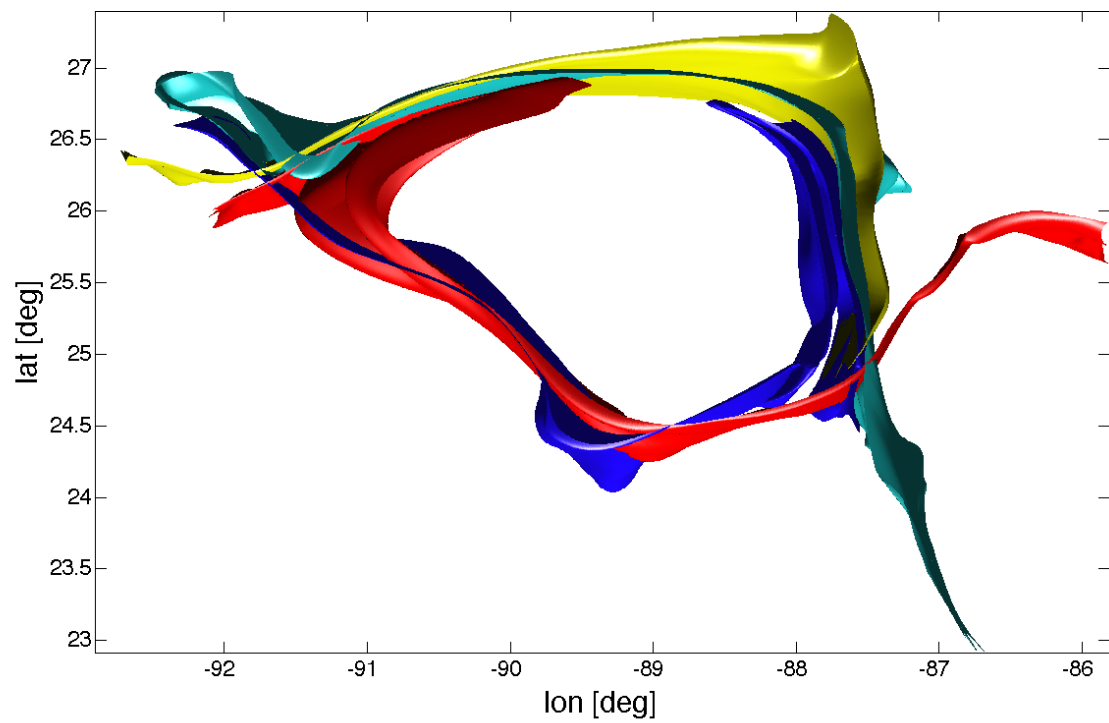
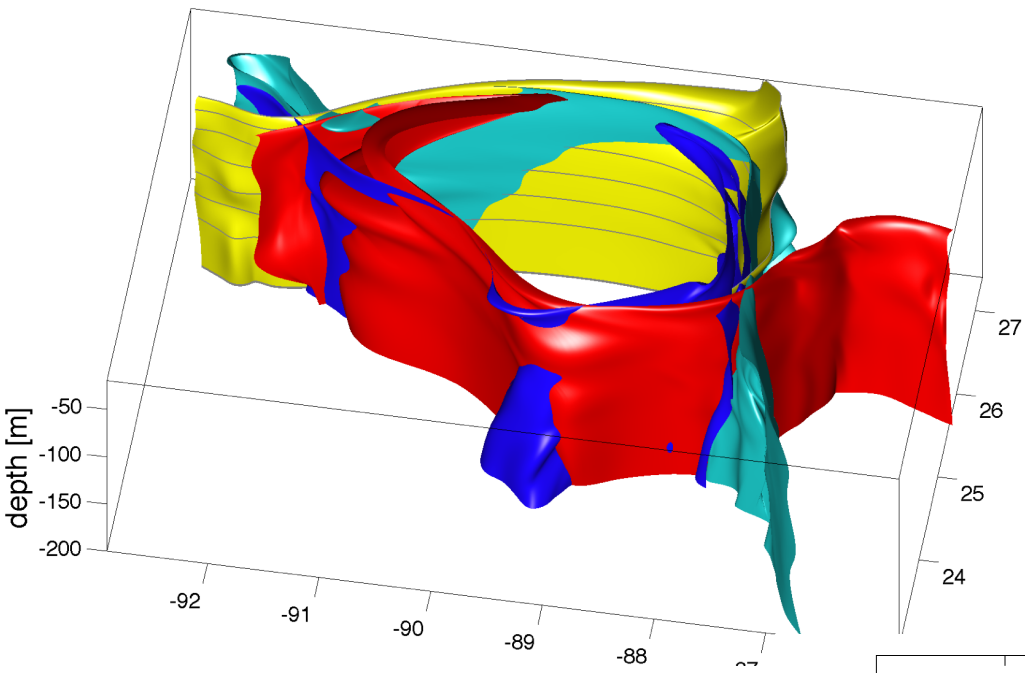


How do manifolds compare with FTLEs ?

**DHTs, manifolds, and lobes near the eddy**

02-Dec-1999 00:00:00





# B&K Conclude

- Material surfaces stitched from 2D analysis revealed coherent lobes with depth
- Material surfaces drop nearly vertically. No evidence of eddy lens structure
- Net inflow at bottom, outflow at top

*Realistic, or artifacts of stitching and/or data assimilation?*

# Feasibility of 3D LCS from 2D Velocities

- Options
  - Follow B & K paradigm
  - Extend FTLE calculations to include vertical shear of horizontal velocities
  - Calculate 3D trajectories using diagnostic vertical velocity
- Strategy
  - Test options with toy models to control vertical velocity and vertical gradients
  - Apply to data-assimilating OGCMs

# Incompressible Models

$$\mathbf{v} = \nabla \times [-\Psi \mathbf{k} + \nabla \times (\Phi \mathbf{k})]$$

$$u = -\frac{\partial \Psi}{\partial y} + \frac{\partial^2 \Phi}{\partial z \partial x}$$

$$v = \frac{\partial \Psi}{\partial x} + \frac{\partial^2 \Phi}{\partial z \partial y}$$

$$w = -\nabla_h^2 \Psi$$

# ABC & Quadrupole Flows

## ABC

$$\Psi = - [C \sin (y + f (t)) + B \cos (x + f (t))]$$

$$\Phi = A [-x \cos (z + f (t)) + y \sin (z + f (t))] - \Psi$$

## Quadrupole

$$\Psi = A (z, t) \sin (\pi x / L_x) \sin (\pi y / L_y)$$

$$\Phi = B (z, t) \cos (\pi x / L_x) \cos (\pi y / L_y)$$



# Strain Tensor and Velocity Gradient

Consider

$$\frac{d\mathbf{x}}{dt} = \mathbf{v}(\mathbf{x}, t), \mathbf{x} \in \Omega \subset \mathcal{R}^3$$

With Solution

$$\mathbf{x} = \mathbf{x}_0 + \int_{t_0}^t \mathbf{v}(\mathbf{x}, \tau) d\tau$$

And Strain Tensor

$$\frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} = \mathbf{I} + \int_{t_0}^t \frac{\partial \mathbf{v}}{\partial \mathbf{x}} \cdot \frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} d\tau$$

# Cauchy-Green Tensor and FTLE

Cauchy-Green

$$\mathbf{C} = \left( \frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} \right)^T \left( \frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} \right)$$

FTLE

$$\Lambda(t; t_0, \mathbf{x}_0) = \frac{\log \left( \sqrt{\lambda_{max}(\mathbf{C})} \right)}{t - t_0}$$

# Strain Tensor

3D

$$\begin{bmatrix} \frac{\partial x}{\partial x_0} & \frac{\partial x}{\partial y_0} & \frac{\partial x}{\partial z_0} \\ \frac{\partial y}{\partial x_0} & \frac{\partial y}{\partial y_0} & \frac{\partial y}{\partial z_0} \\ \frac{\partial z}{\partial x_0} & \frac{\partial z}{\partial y_0} & \frac{\partial z}{\partial z_0} \end{bmatrix}$$

2D

$$\begin{bmatrix} \frac{\partial x}{\partial x_0} & \frac{\partial x}{\partial y_0} \\ \frac{\partial y}{\partial x_0} & \frac{\partial y}{\partial y_0} \end{bmatrix} \rightsquigarrow \frac{\partial}{\partial z_0} \begin{bmatrix} \mathbf{x} \\ \mathbf{v} \end{bmatrix} = 0$$

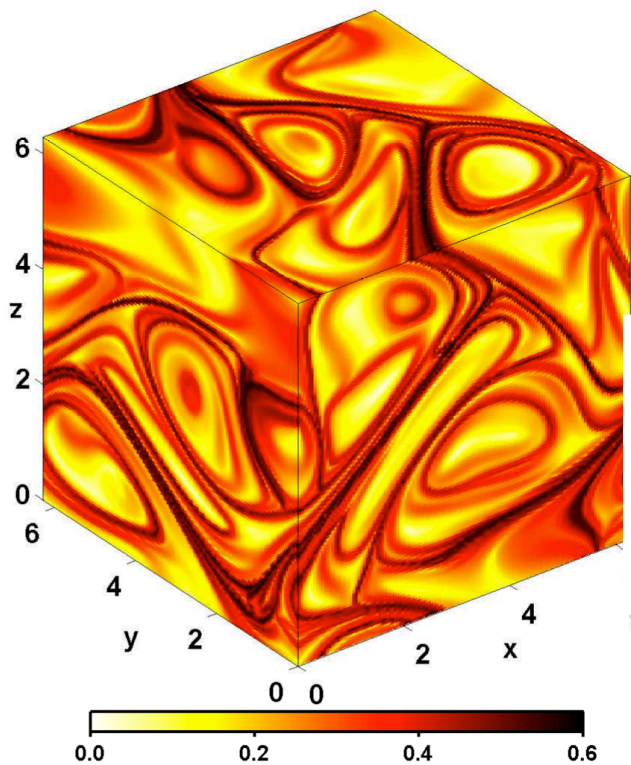
$$\nabla_h z = 0$$

3D2D

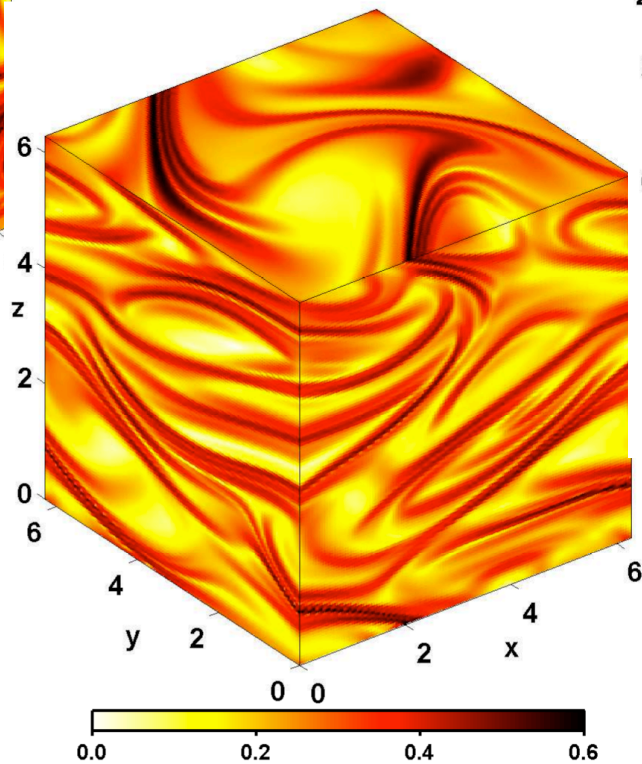
$$\begin{bmatrix} \frac{\partial x}{\partial x_0} & \frac{\partial x}{\partial y_0} & \frac{\partial x}{\partial z_0} \\ \frac{\partial y}{\partial x_0} & \frac{\partial y}{\partial y_0} & \frac{\partial y}{\partial z_0} \\ 0 & 0 & 1 \end{bmatrix}$$

$$\rightsquigarrow \frac{\partial}{\partial x_0, \partial y_0} \begin{bmatrix} z \\ w \end{bmatrix} \sim 0$$

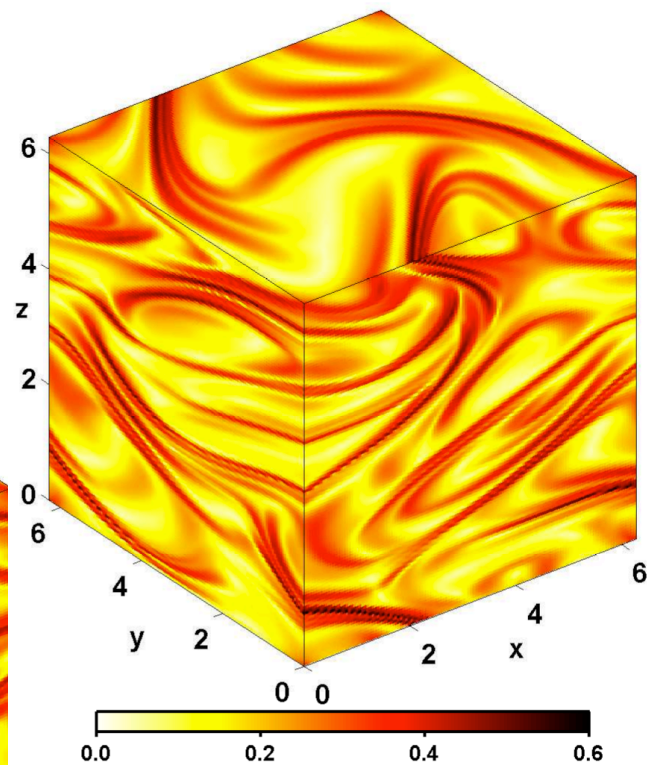
# ABC Flow



3D

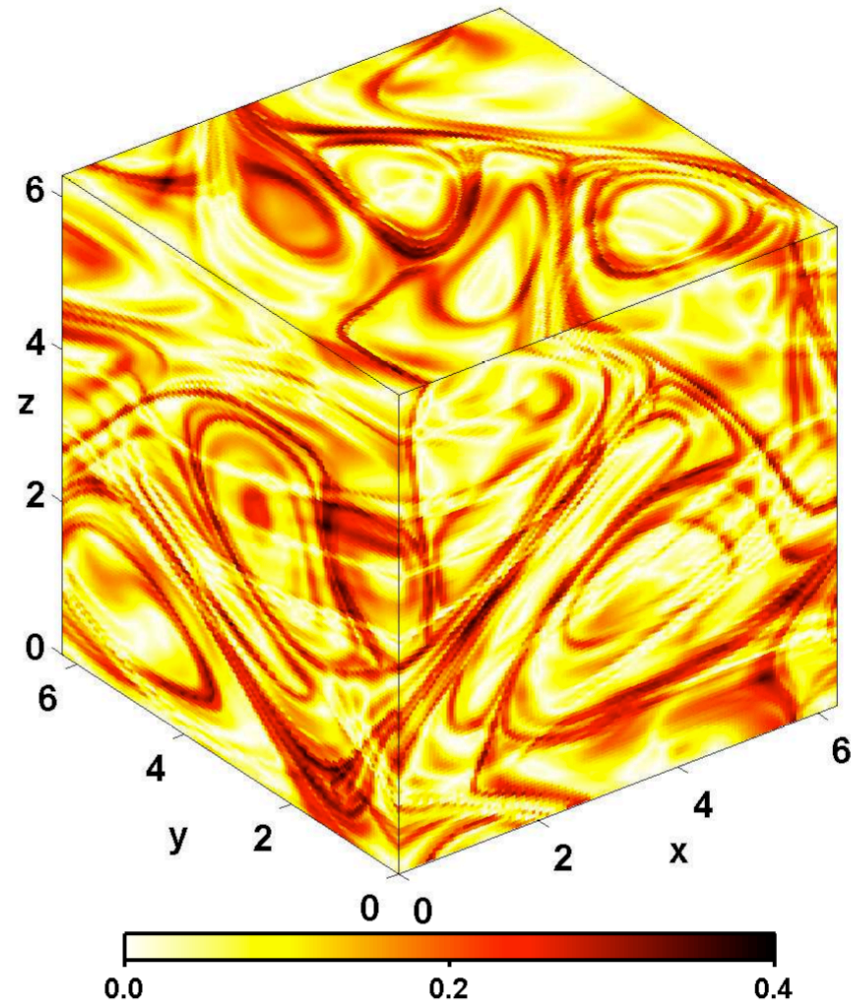
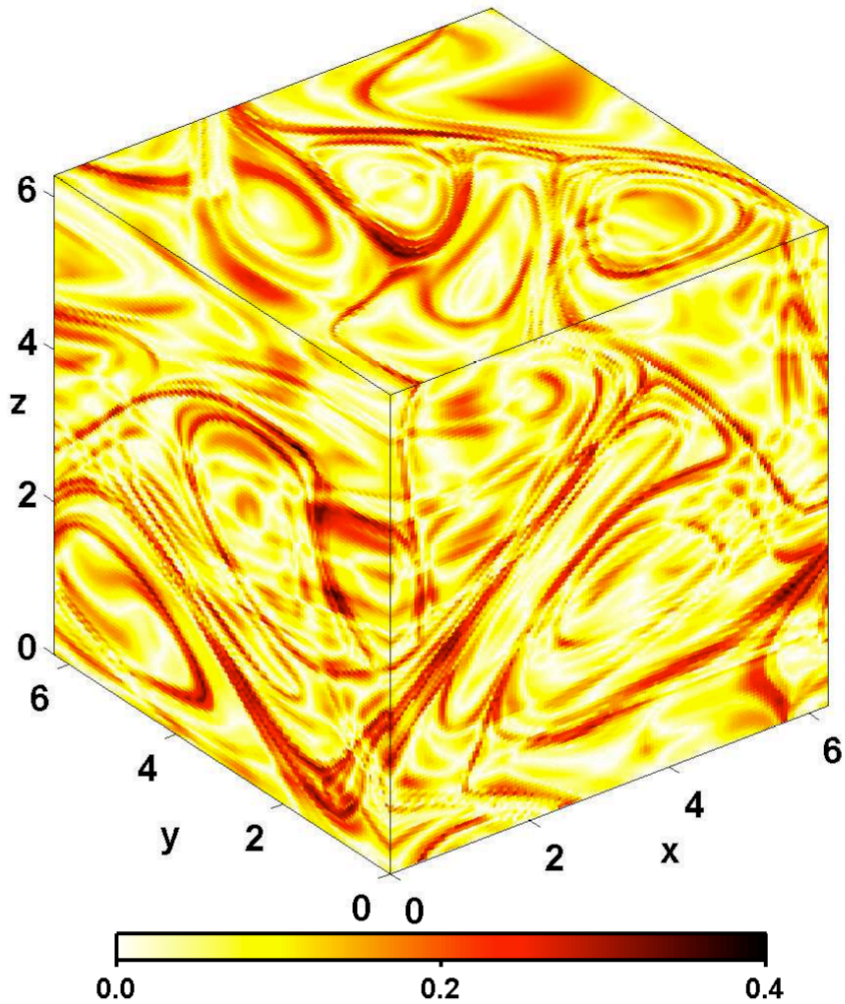


3D2D



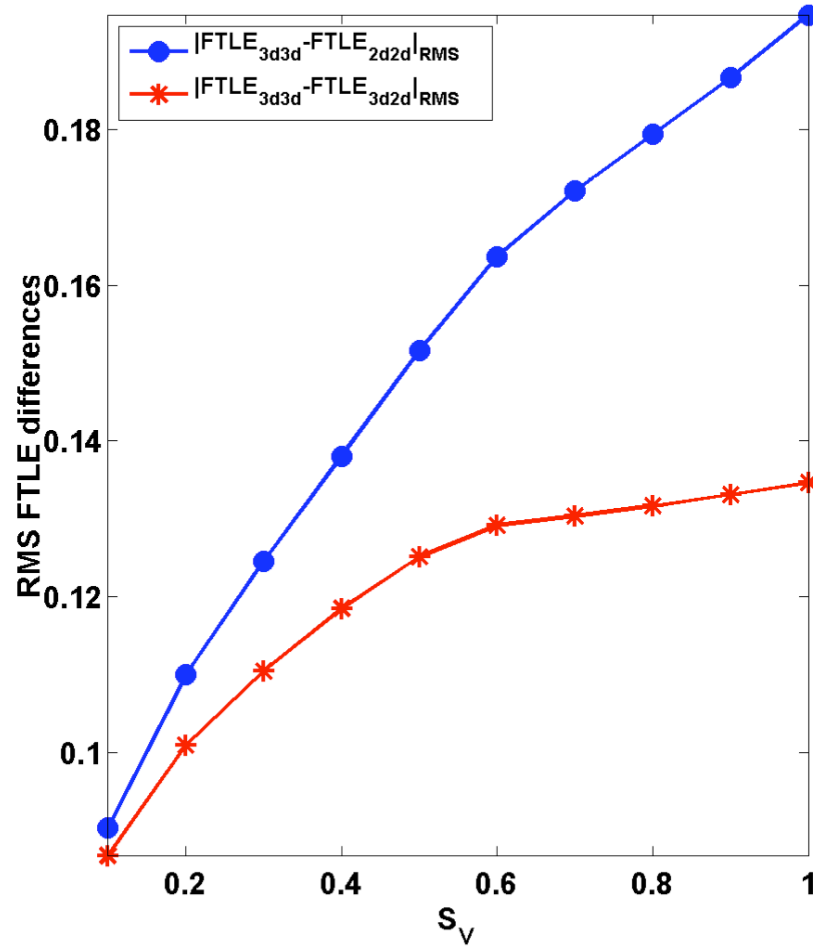
2D

# FTLE Differences



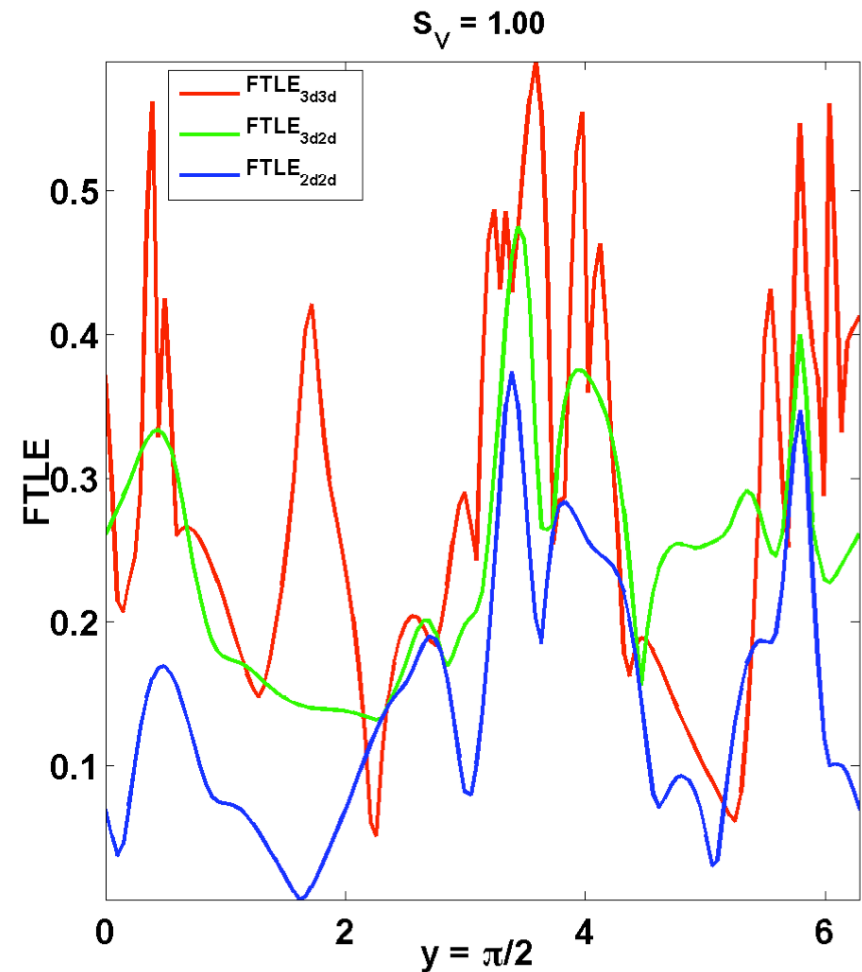
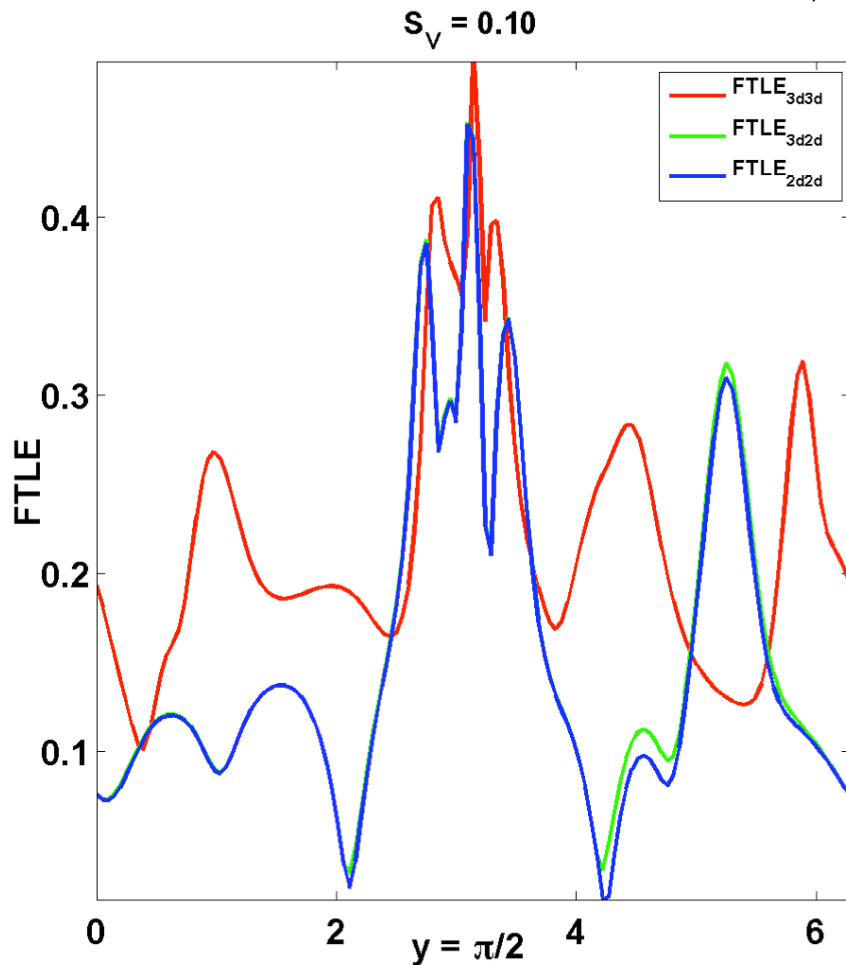
# RMS FTLE Differences

$$S_V = \sqrt{\left(\frac{\partial u}{\partial z}\right)^2 + \left(\frac{\partial v}{\partial z}\right)^2}$$



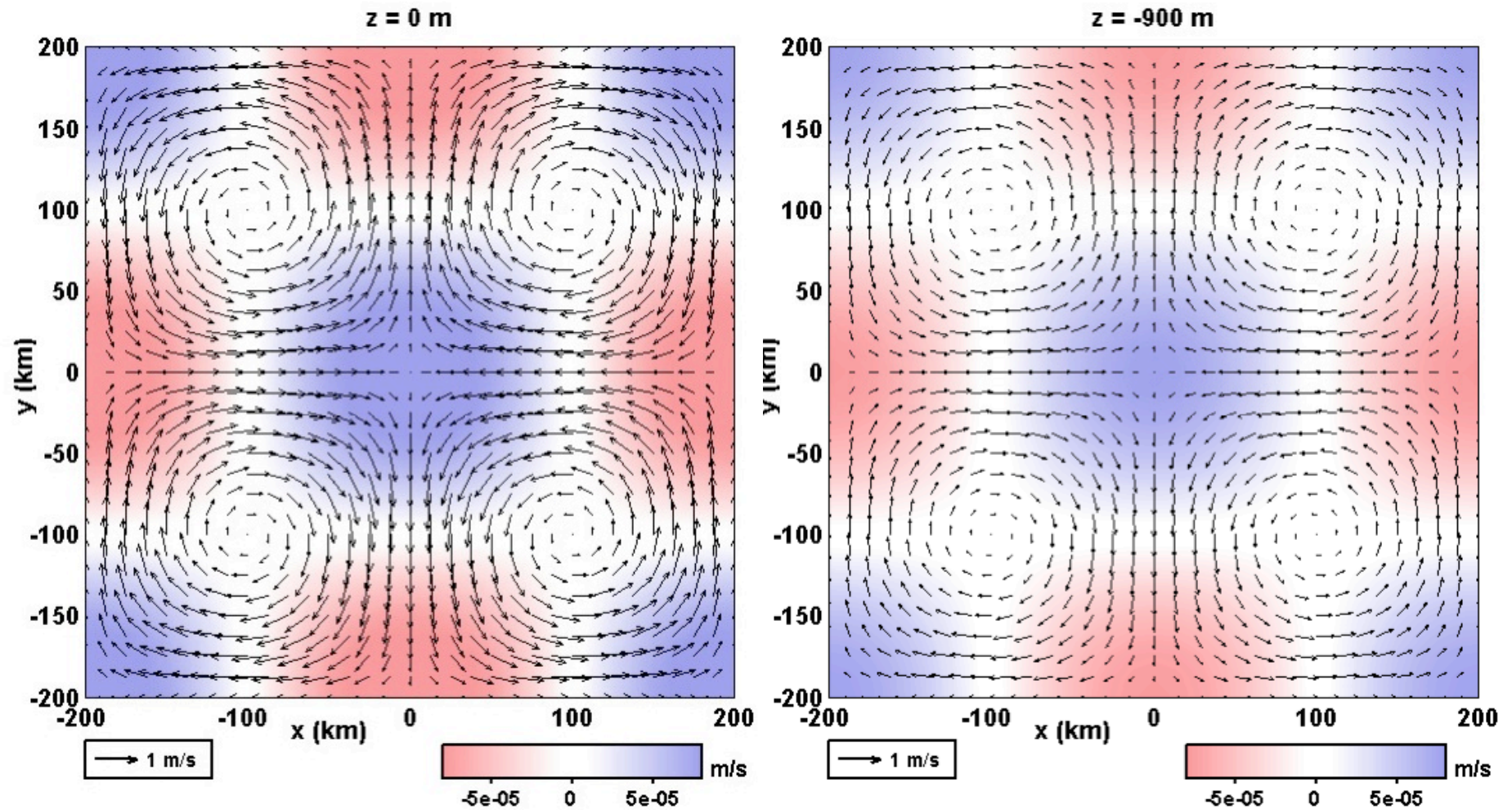
# It Doesn't Look Good for Reduced Representations of Cauchy - Green

$$S_V = \sqrt{\left(\frac{\partial u}{\partial z}\right)^2 + \left(\frac{\partial v}{\partial z}\right)^2}$$



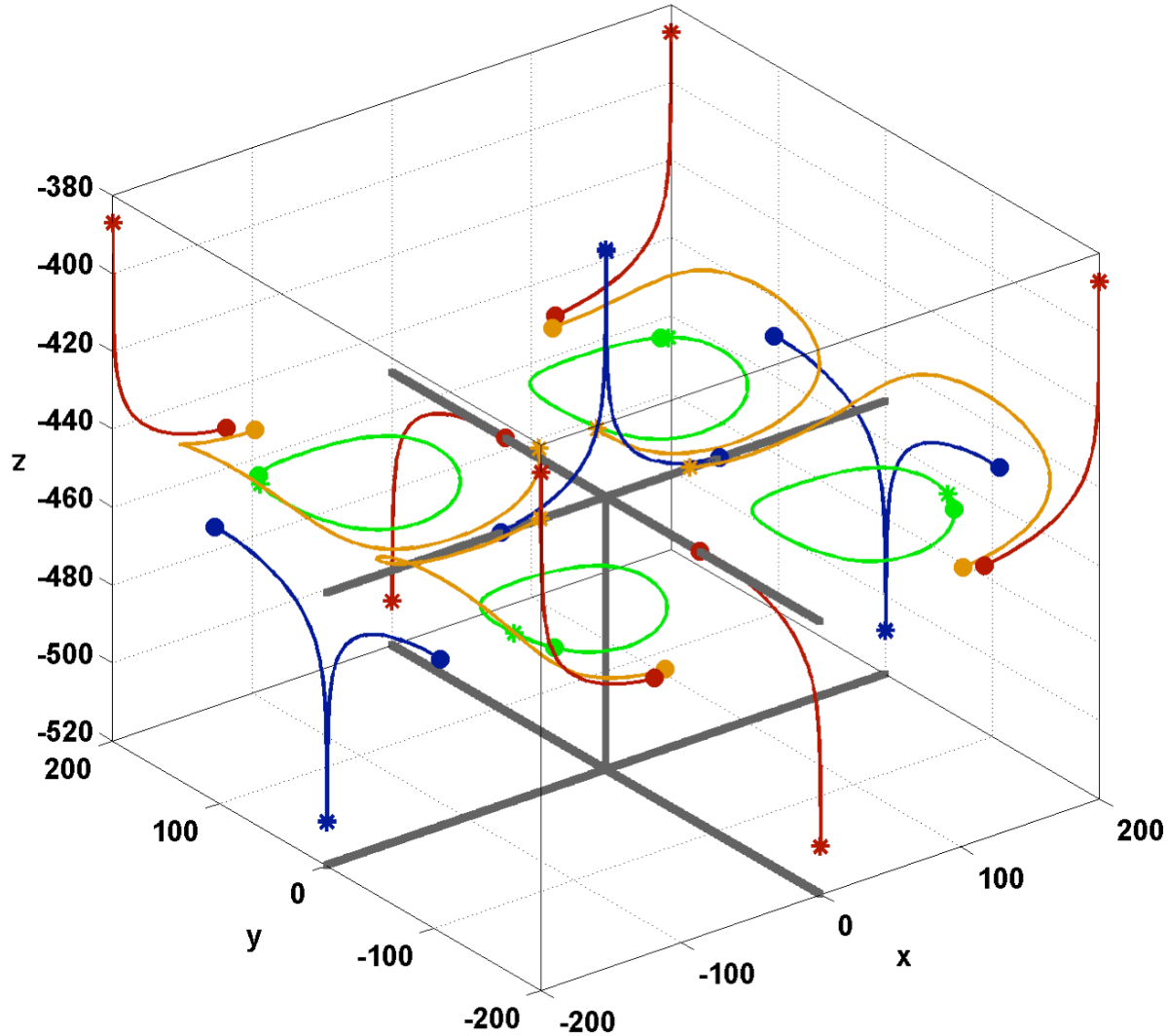


# Quadrupole - Eulerian View

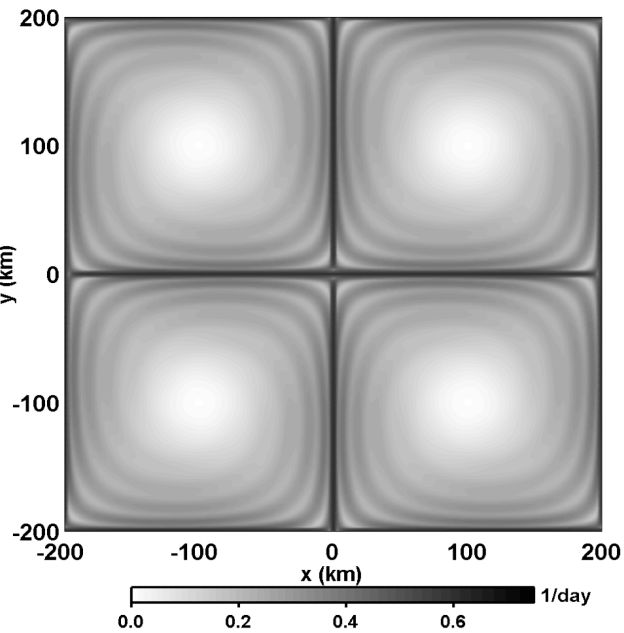




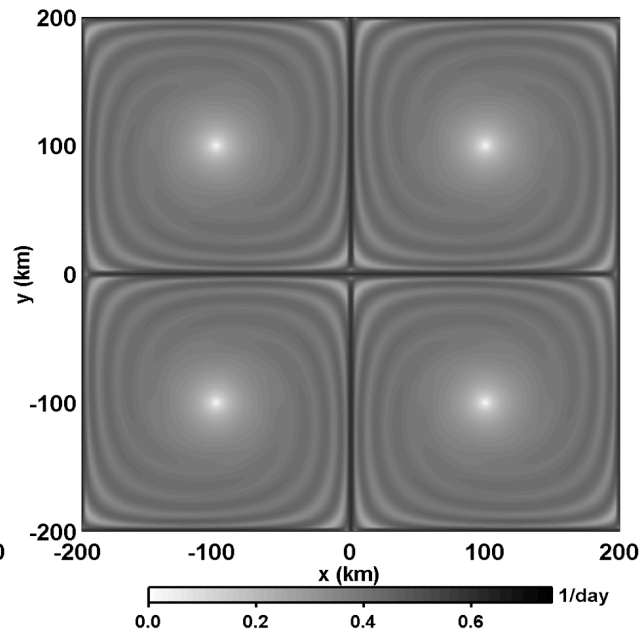
# Lagrangian View at - 450m



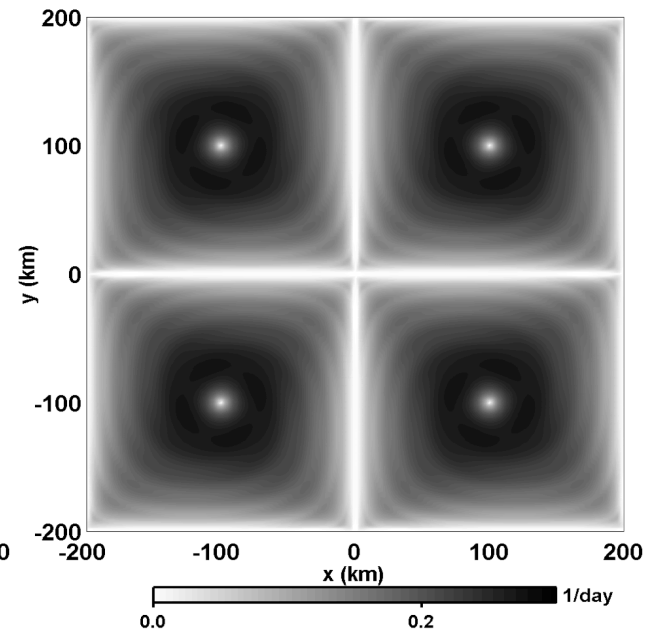
# FTLE Comparison – 450m



3D



3D2D

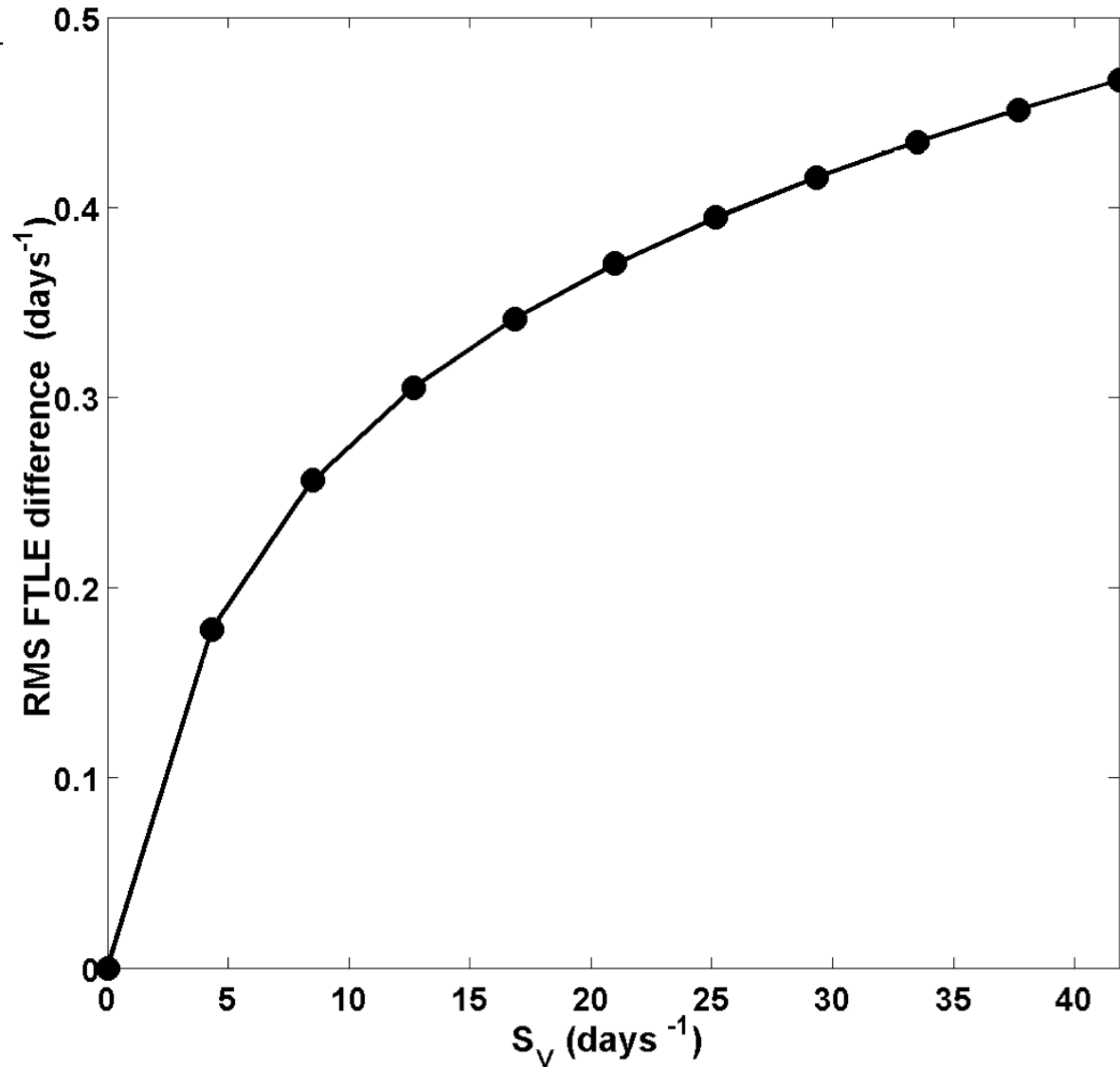


Difference

# Effect of Vertical Shear on FTLE

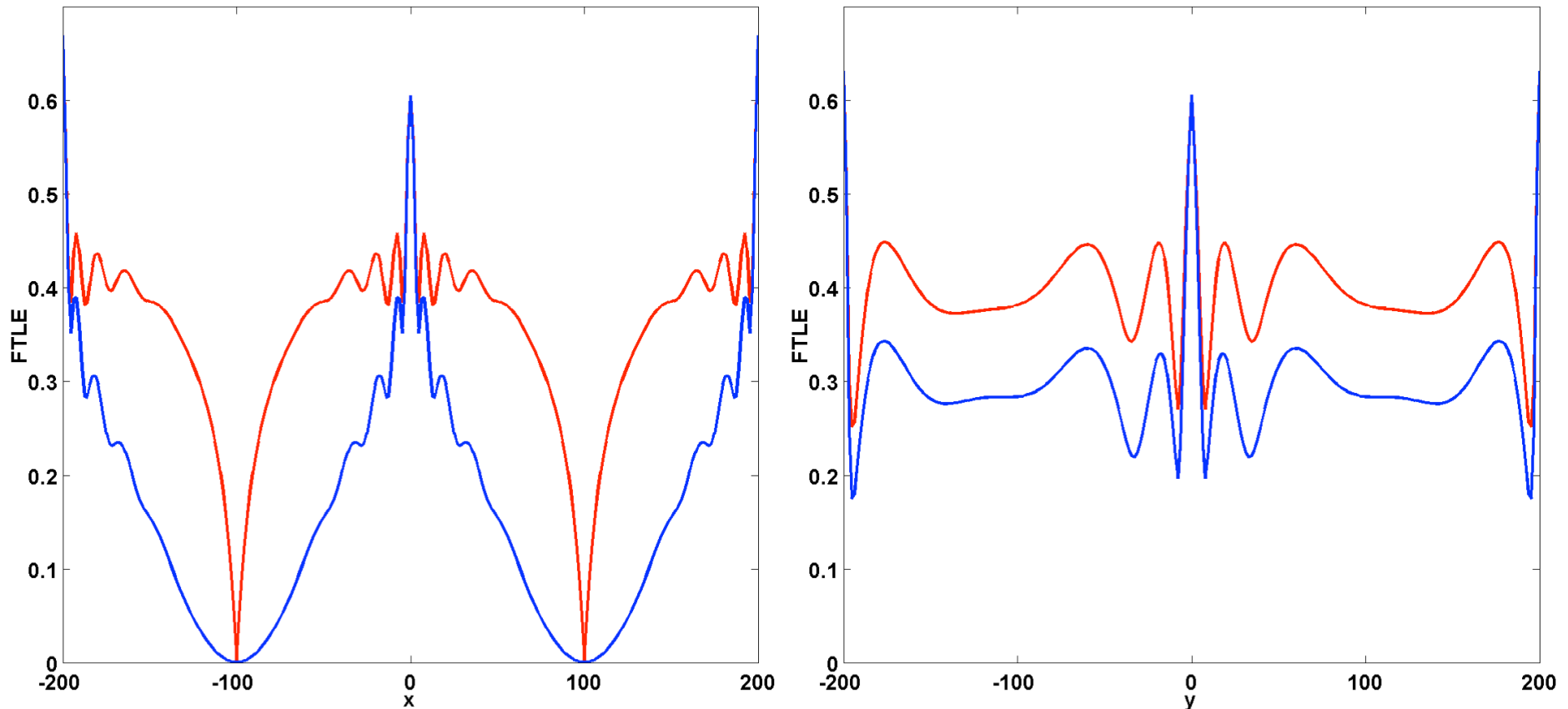
$$S_V = \sqrt{\left(\frac{\partial u}{\partial z}\right)^2 + \left(\frac{\partial v}{\partial z}\right)^2}$$

$$S_V \leq 2N$$



# FTLE Comparison

Figure 4. Profiles of FTLE3d (red) and FTLE2d (blue) at  $z = -450$  m. Left: Profile along  $y = -100$  km; right: Profile along  $x = -12.5$  km.



# What to Expect in Ocean

For large FTLE  $\partial x \approx \partial y = \Delta_H \approx 50$  km and  $\partial z = \Delta_V \approx 0.4$ km

And  $\partial x_0 \approx \partial y_0 = \Delta_{H_0} \approx 5$ km and  $\partial z_0 \approx \Delta_{V_0} \approx 0.03$ km

Then

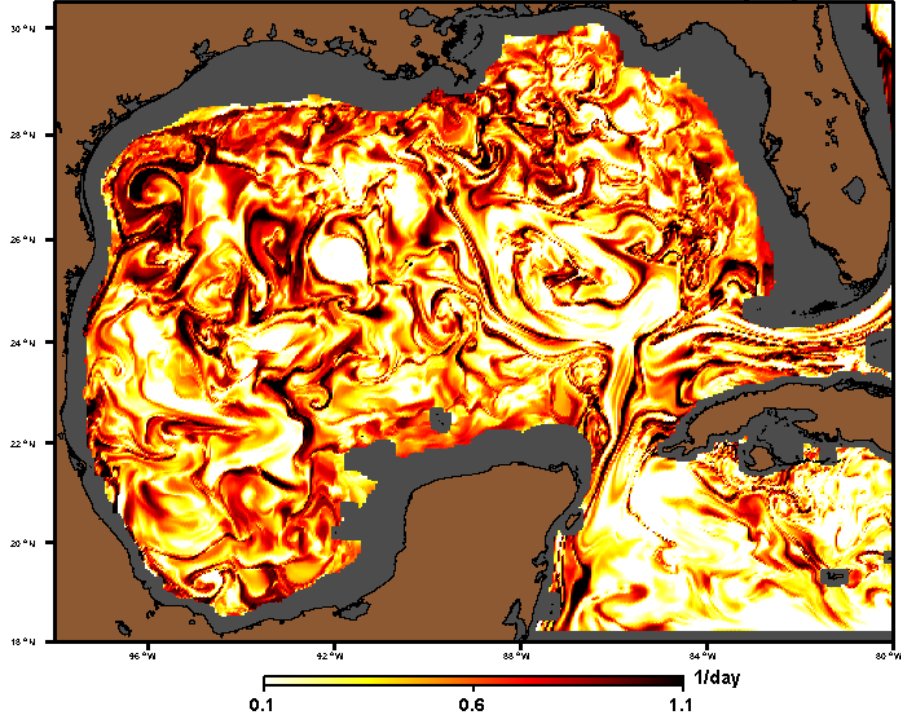
$$\left[ \begin{array}{ccc} \frac{\partial x}{\partial x_0} = a \approx 10 & \frac{\partial x}{\partial y_0} = b \approx 10 & \frac{\partial x}{\partial z_0} = \Gamma \approx 1.7 \cdot 10^3 \\ \frac{\partial y}{\partial x_0} = c \approx 10 & \frac{\partial y}{\partial y_0} = d \approx 10 & \frac{\partial y}{\partial z_0} = \Gamma \approx 1.7 \cdot 10^3 \\ \frac{\partial z}{\partial x_0} = \epsilon \approx 4 \cdot 10^{-2} & \frac{\partial z}{\partial y_0} = \epsilon \approx 4 \cdot 10^{-2} & \frac{\partial z}{\partial z_0} = e \approx 1.4 \cdot 10. \end{array} \right]$$

And

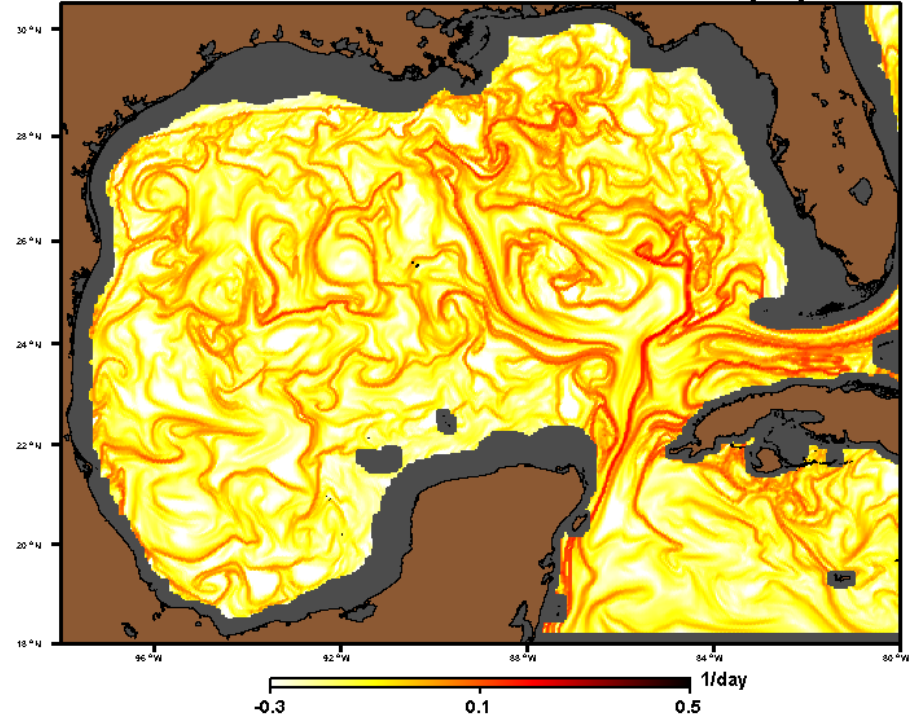
$$\left( \frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} \right)^T \left( \frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} \right) = \left[ \begin{array}{ccc} a^2 + c^2 & ab + cd & \Gamma(a + b) \\ ab + cd & b^2 + d^2 & \Gamma(b + d) \\ \Gamma(a + c) & \Gamma(b + d) & 2\Gamma^2 + e^2 \end{array} \right]$$

# Application to GoM HYCOM

06/15/2010 GOM HYCOM 3D FTLE at 25.0m 10 day trajectories



06/15/2010 GOM HYCOM 2D FTLE at 25.0m 10 day trajectories



# What We Get

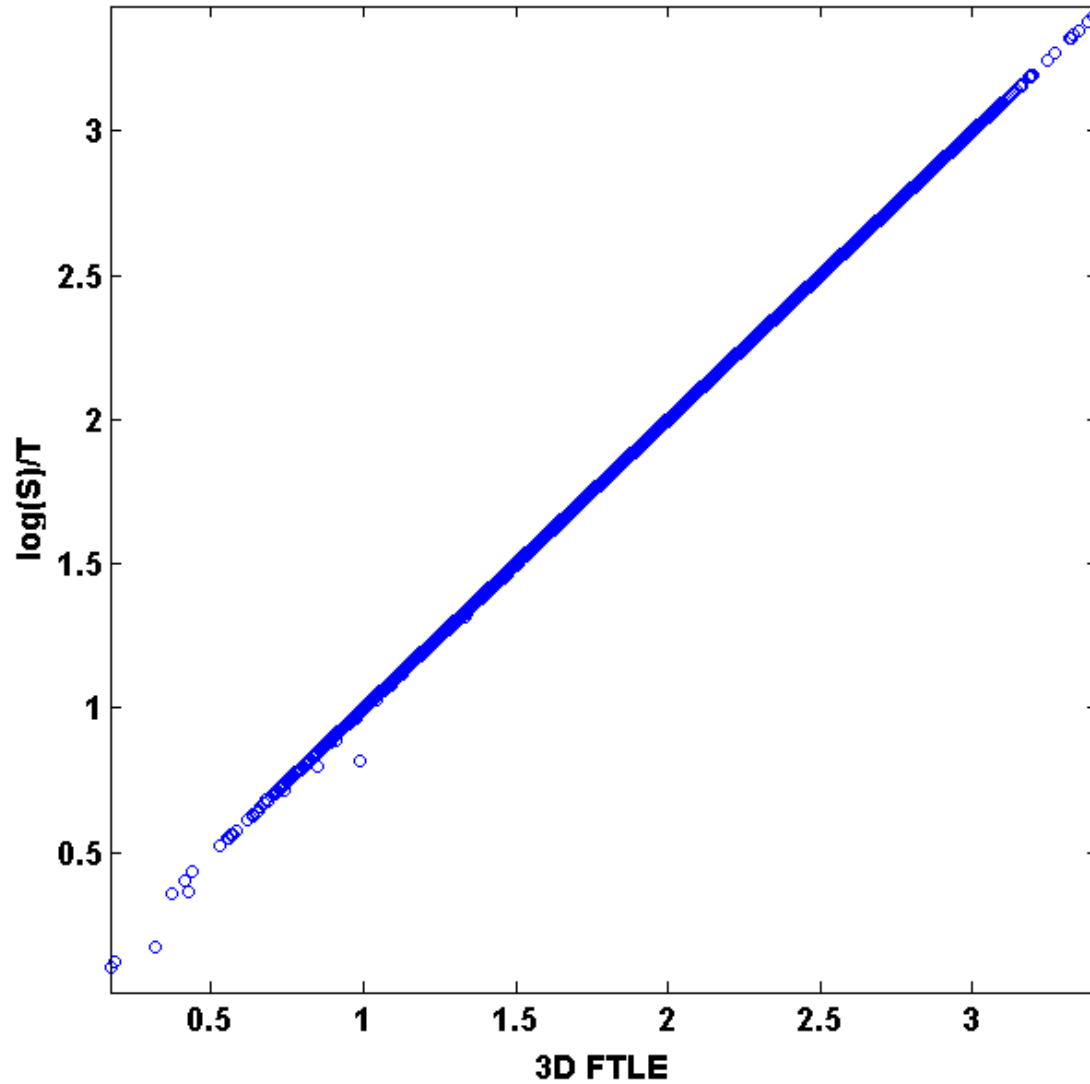
$$\frac{\partial \mathbf{x}}{\partial \mathbf{x}_0} = \begin{bmatrix} 42 & 39 & 4057 \\ 21 & 17 & 1753 \\ 0 & 0 & 1 \end{bmatrix}$$

These matrices borderline ill-conditioned.

But only need largest eigenvalue, which is nearly:

$$C_{33} = 1 + \left( \frac{\partial x}{\partial z_0} \right)^2 + \left( \frac{\partial y}{\partial z_0} \right)^2 = 6.5 \cdot 10^4$$

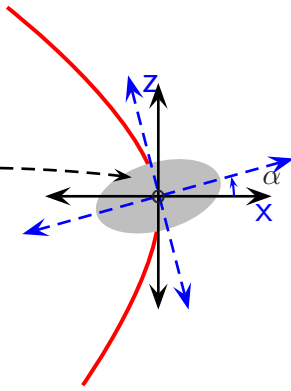
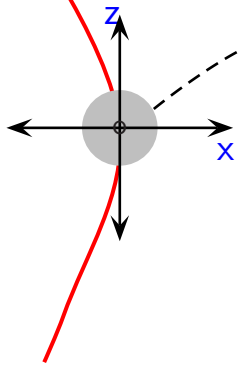
# 3D2D FTLE vs $C_{33}$



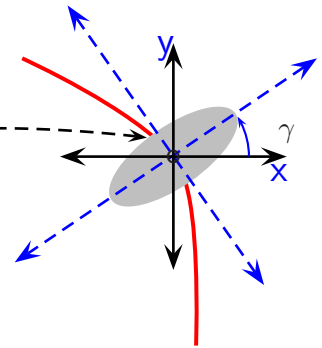
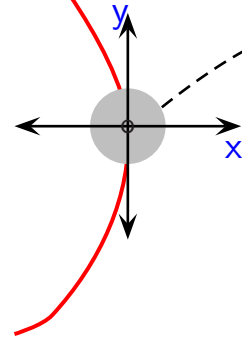


# 3D2D Eigendirections

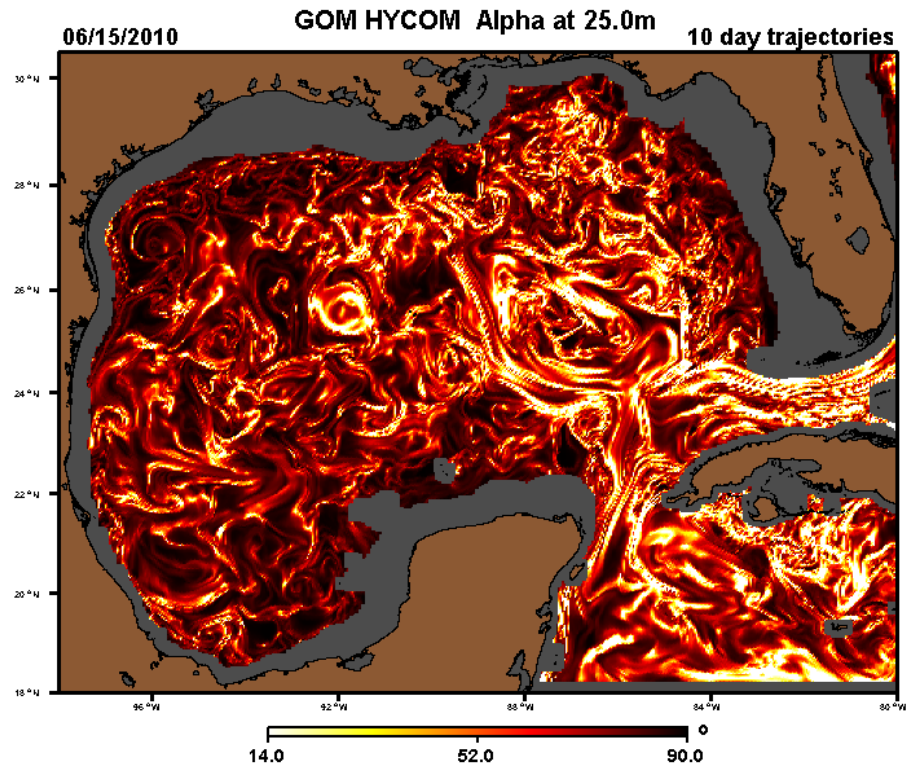
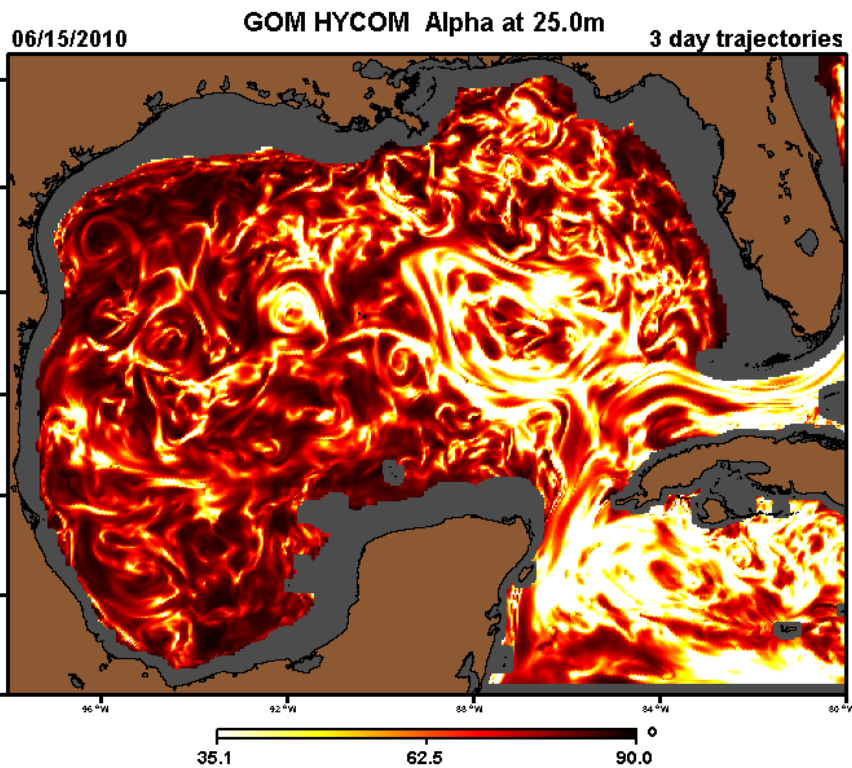
repelling material line



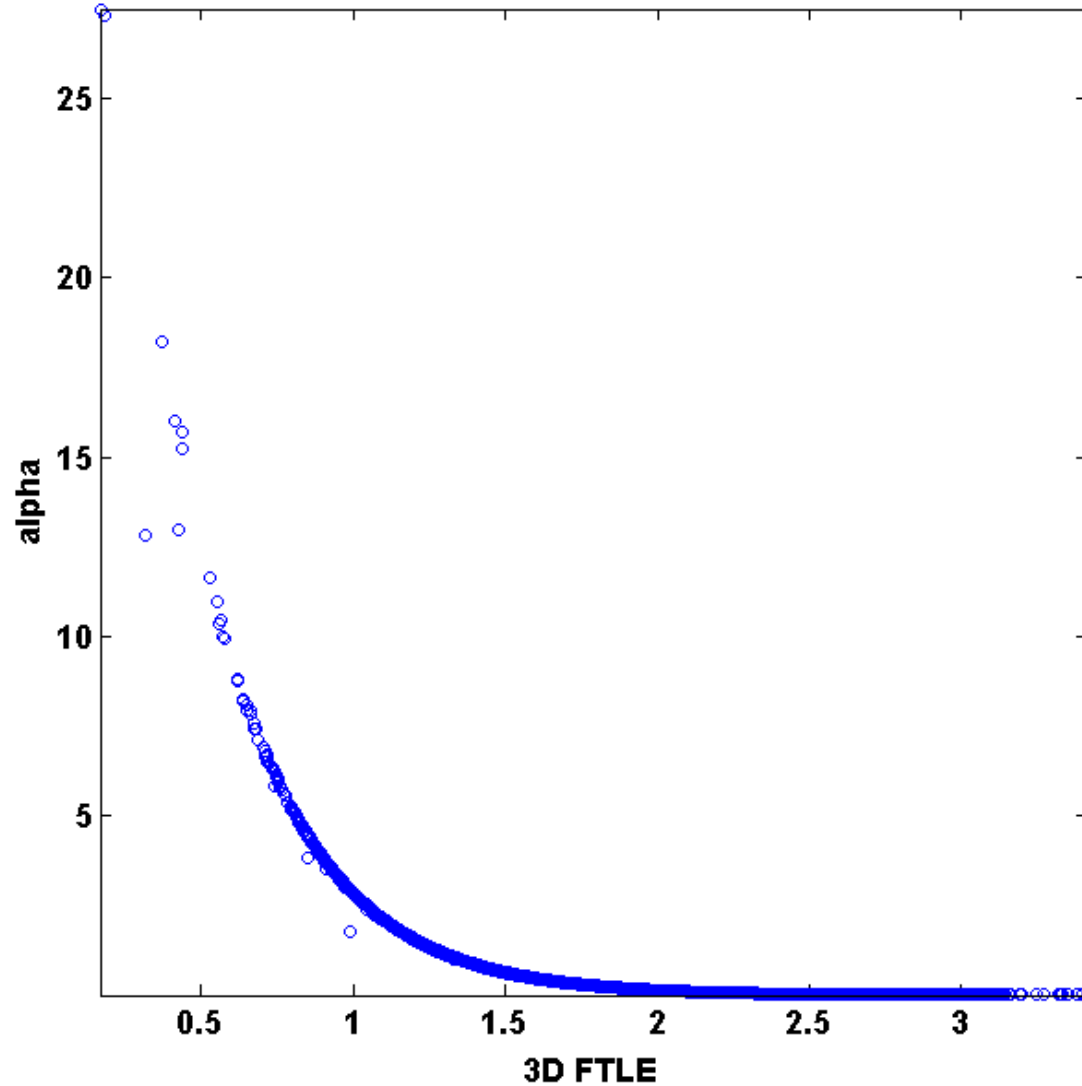
repelling material line



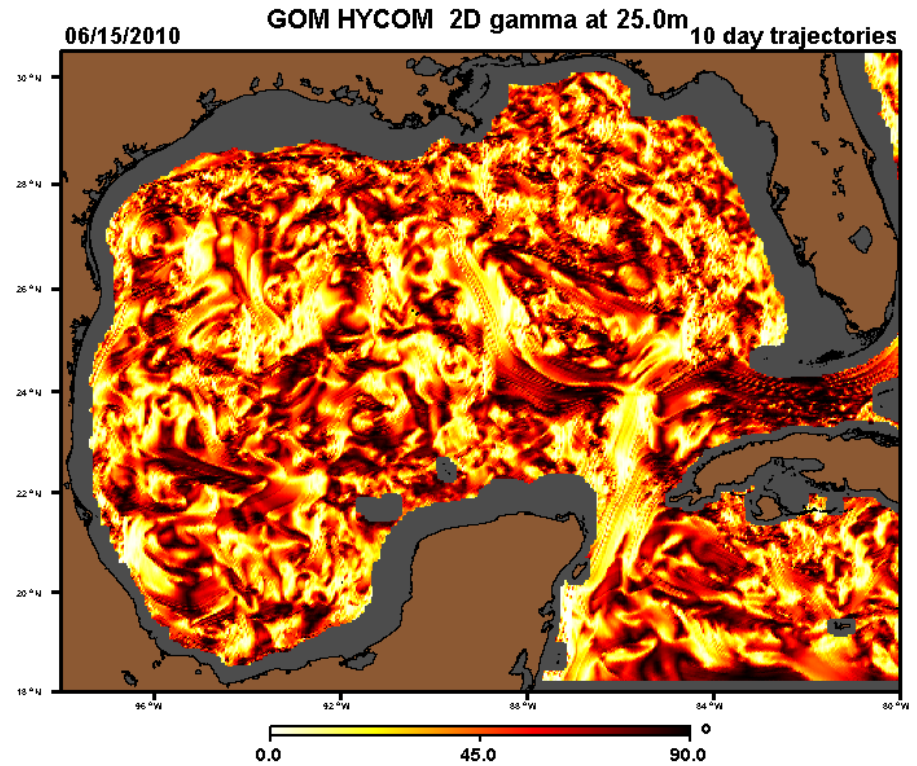
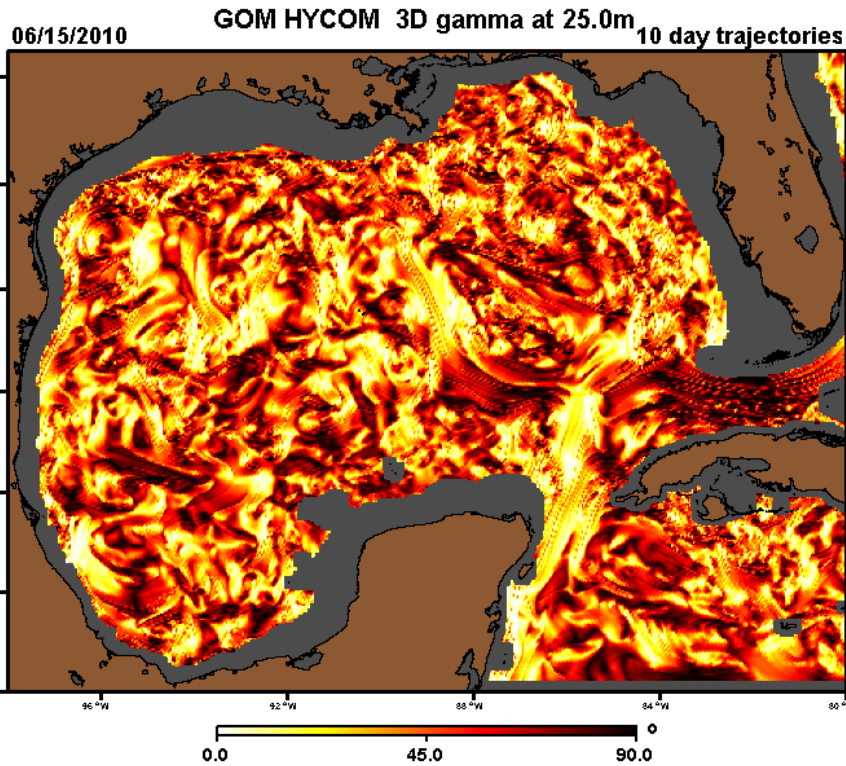
# Vertical Angle of Max Stretch



# 3D2D FTLE Directions



# Horizontal Angle of Max Stretch



# Discussion

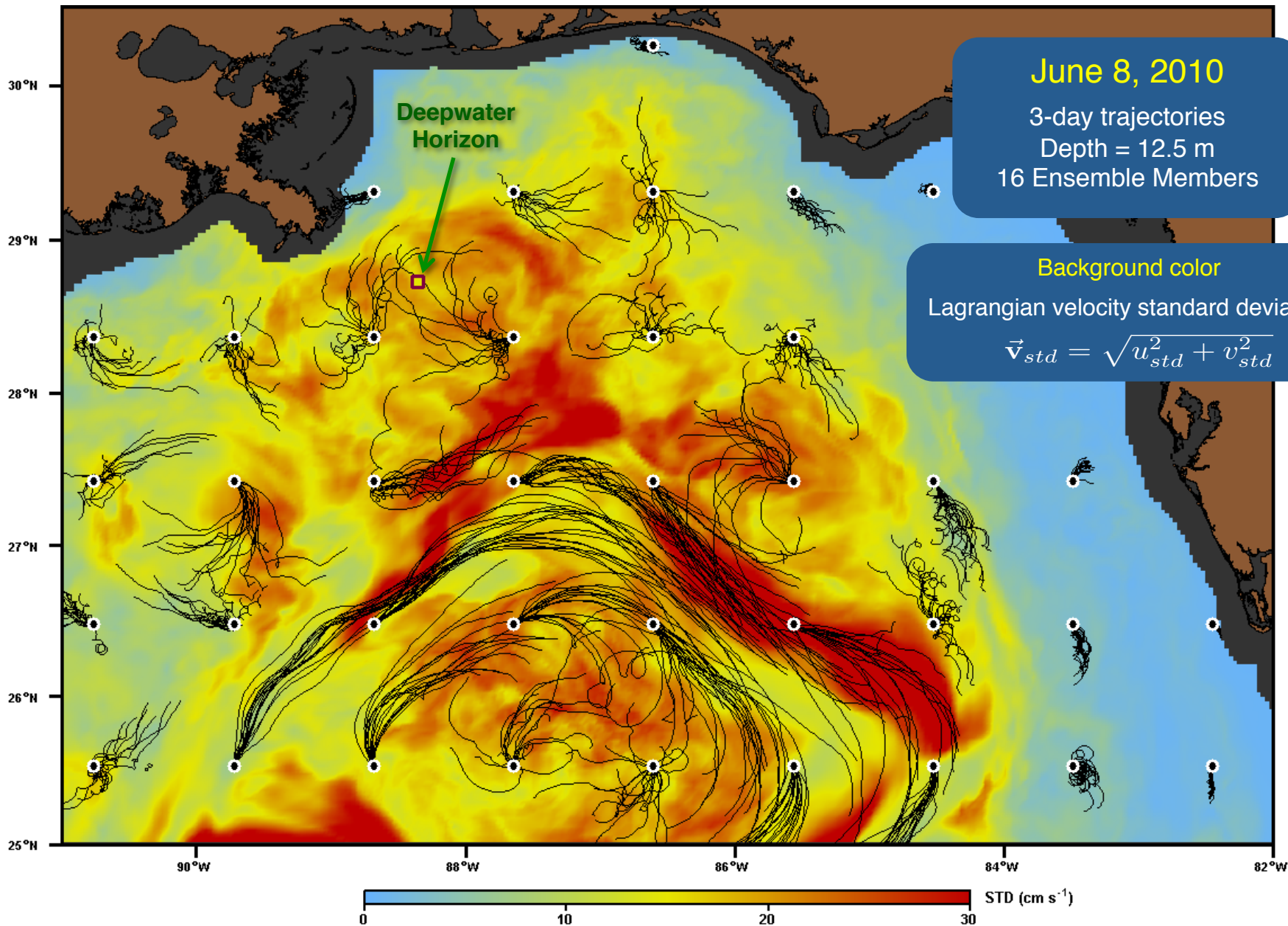
- In the ocean
  - 2D velocity fields capture essence of 3D FTLE fields
  - Large stretch is nearly horizontal
- In GoM 3D FTLE Surrogate:  $S_V = (\partial x / \partial z_0)^2 + (\partial y / \partial z_0)^2$
- Next
  - Test  $S_V$  with other metrics
  - Construct 2D transport barriers from data-assimilating OGCMs
  - **No more Mr Nice Guy on 2D pictures!**
  - Eddy formation, census, and 3D transport

# Lagrangian Predictability

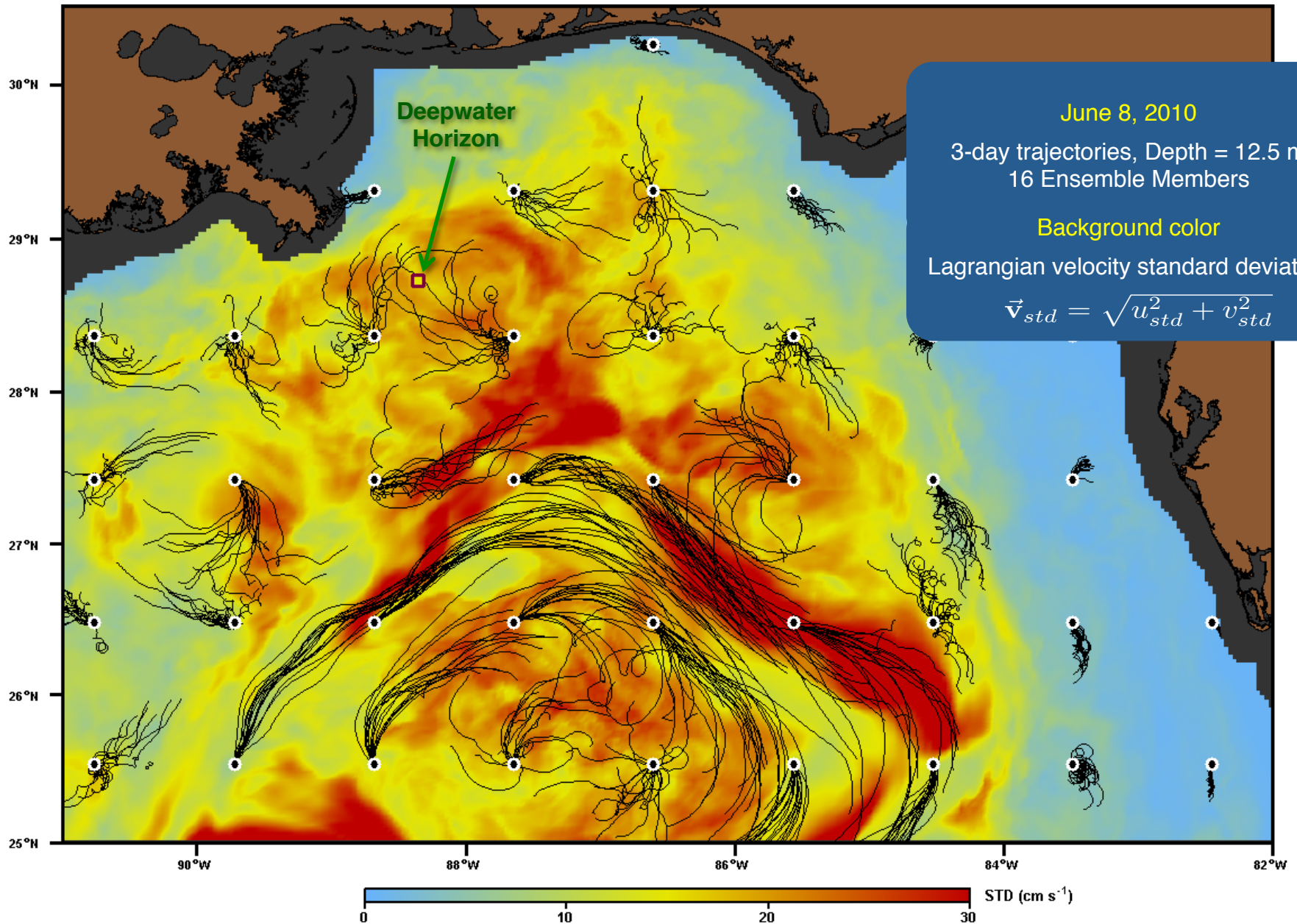
## Ensembles



# RELO in the Northern GoM: Ensemble Spread

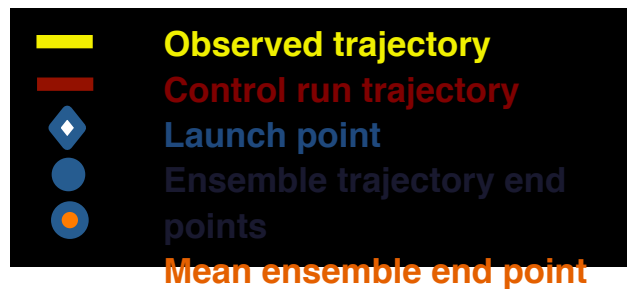
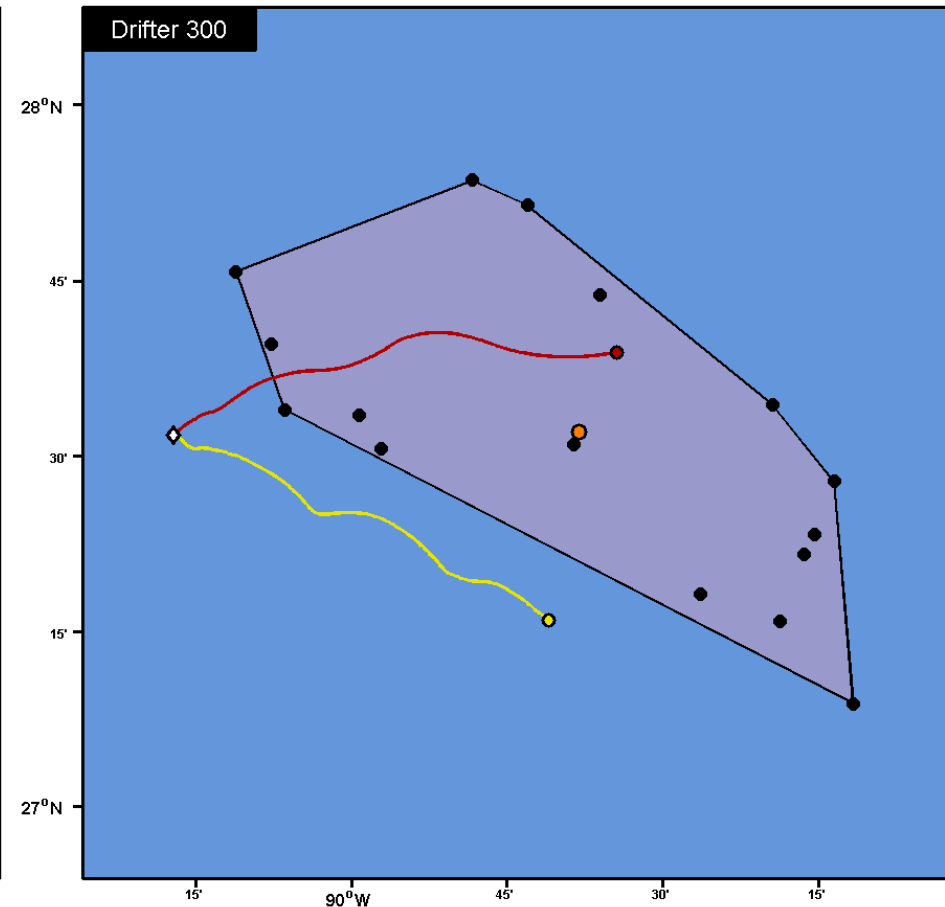
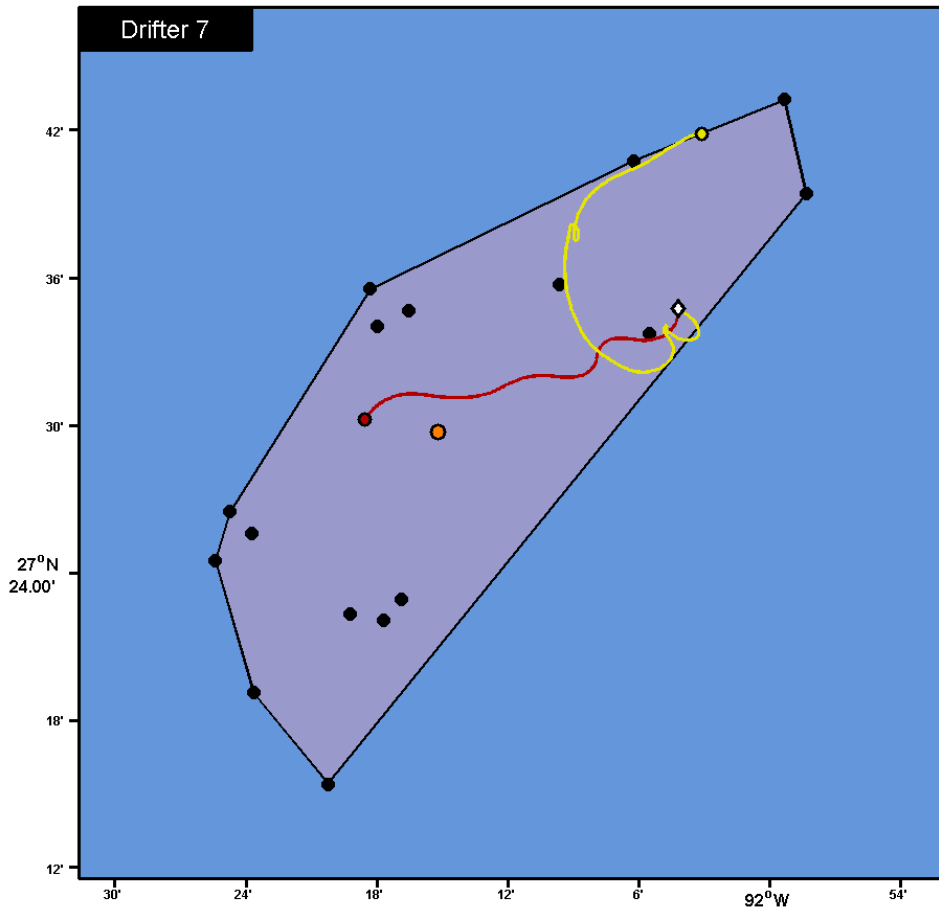


# RELO in the Northern GoM: Ensemble Spread

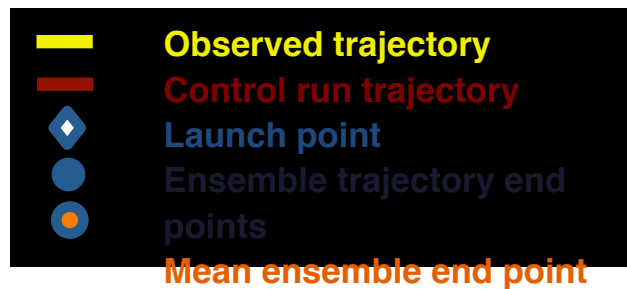
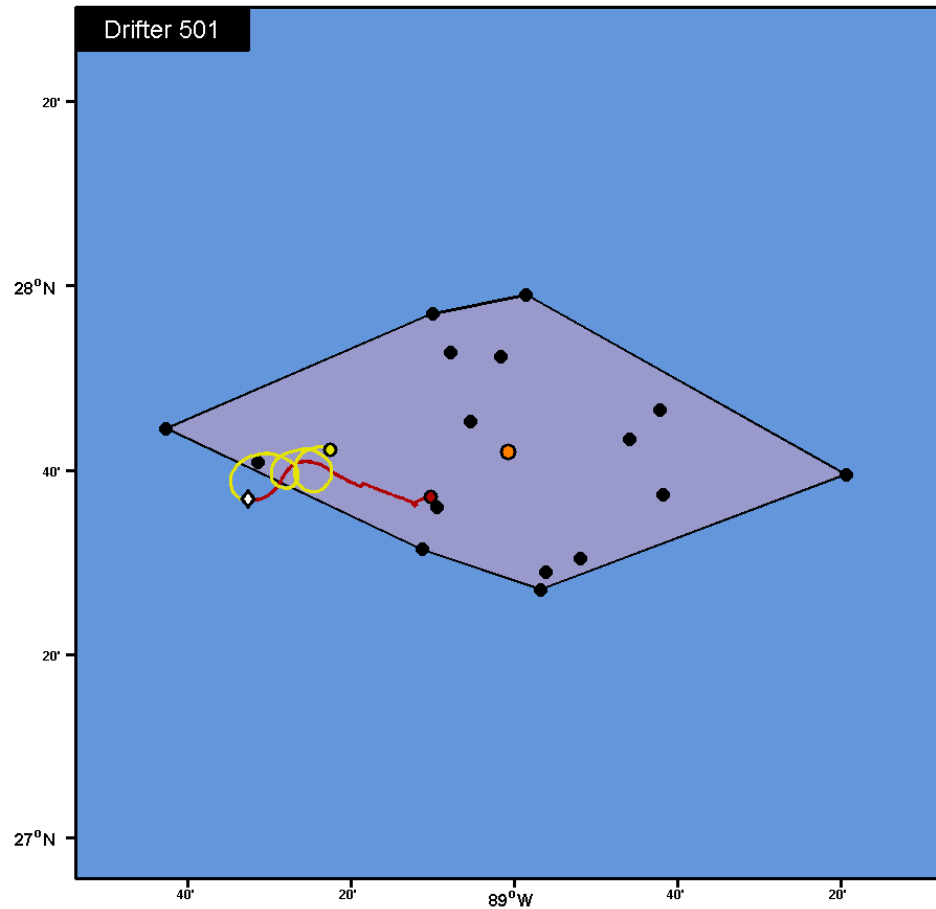
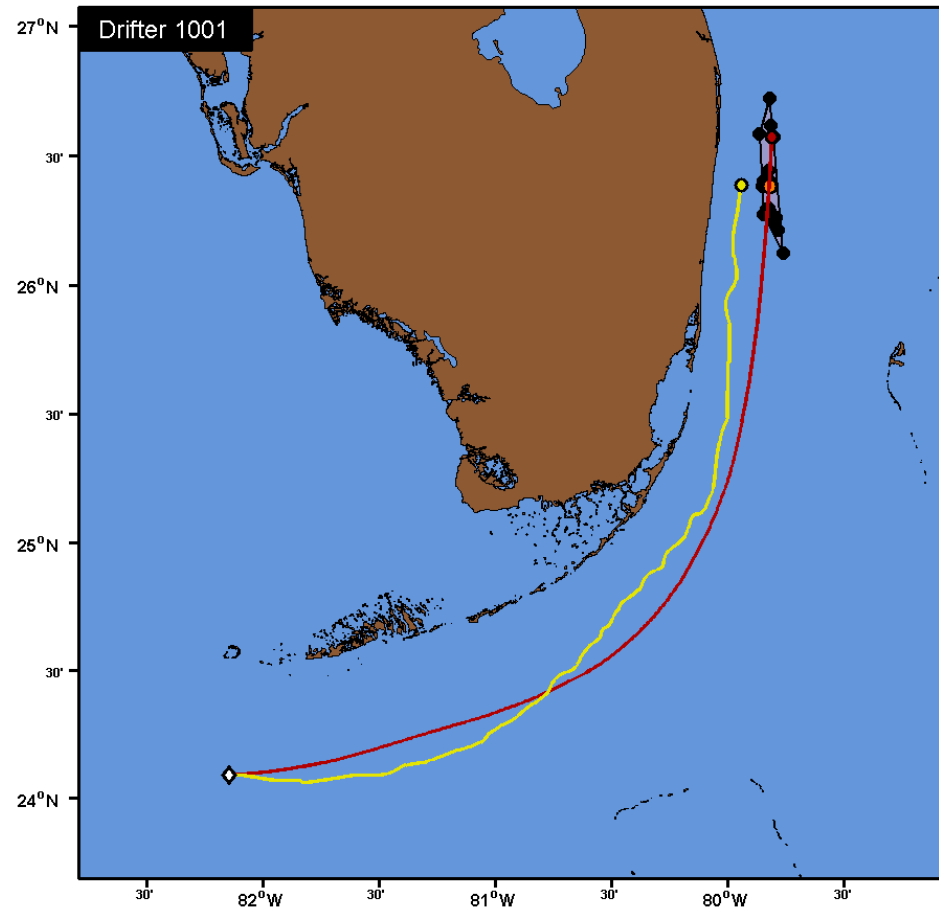




# Example RELO Ensemble



# Example RELO Ensemble



# All is Based on Data Assimilating Models – How Good are They?

